

Attachment I

Marine Geological Investigations in the Beaufort Sea in 1981 and Preliminary Interpretations for region from the Canning River to the Canadian Border.

By: Erk Reimnitz, Peter W. Barnes, Peter W. Minkler, Douglas M. Rearic,
Edward W. Kempema, and Thomas Reiss.

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INTRODUCTION

The USGS vessel R/V **KARLUK** ran approximately 1000 km of geophysical **tracklines** on the inner shelf of the Beaufort Sea, Alaska from July 14 to August 20, 1981. In addition to the **trackline** surveys, 37 sediment grab samples were collected, one area was investigated by SCUBA divers, and 5 sites were monitored with Ocean Bottom Seismographs (**OBS**), three per site. The R/V **KARLUK** left the Beaufort Sea on August 20 to support investigations by Drs. Ralph Hunter and Larry Phillips in the **Chukchi** Sea.

In our 1981 field efforts, the emphasis was on reconnaissance data collection from the eastern sector, between the Canning River and the international border. This **work** was accomplished in two legs, the first one under **P.W. Barnes**, the second under **Erk Reimnitz**. Ice and weather conditions were about average to favorable for inner shelf navigation during the first half of the available open-water season. In this report we outline the general scope of our 1981 field efforts in the Beaufort Sea, the types of **equipment** used, list much of the data gathered, present those parameters already extracted from the geophysical records, and give preliminary interpretations of our findings.

DESCRIPTION OF FIELD OPERATIONS

Reconnaissance work - Our primary goal, a reconnaissance survey from the Canning River to the Canadian border, where almost no inner **shelf** data is available, was accomplished (see Fig. 1). Geophysical lines were run as far offshore as ice concentrations allowed. All lines from the Canning River eastward extend seaward into very tight pack ice, beyond which further penetration was impossible. Early in the season this tight pack ice was near the coastline. As the season progressed, lines could extend farther seaward. One bay and one lagoon were surveyed along this shore. Thirty-seven grab samples were collected, mainly on the open shelf. For this reconnaissance **work** navigational control is based on radar fixes and dead reckoning. The probable uncertainty in position ranges from 100 or 200 m near shore, to as much as 3 km under dead reckoning on the seaward ends of several tracklines.

Site-specific work - **Between** the Canning River and the **Colville** River, surveys were site specific. Detailed surveys for preparation of side-scan sonar mosaics with bathymetry were run in four small areas, **two** on Stamukhi Shoal, **one** on the 18-m bench seaward of Narwhal Island, and another one on the 18-m bench seaward of Reindeer Island. Detailed bathymetric surveys were run around the "West DockC" and around two artificial gravel islands: Niakuk 3 and B.I?. 37. **Two test** lines from previous years were **re-run** (first run in **1973**, see **Reimnitz**, et al., 1977; and Barnes, et al., 1978) and two new test lines were established with side-scan sonar to determine yearly rates of ice gouging. For all of these detailed surveys, positions were plotted using a

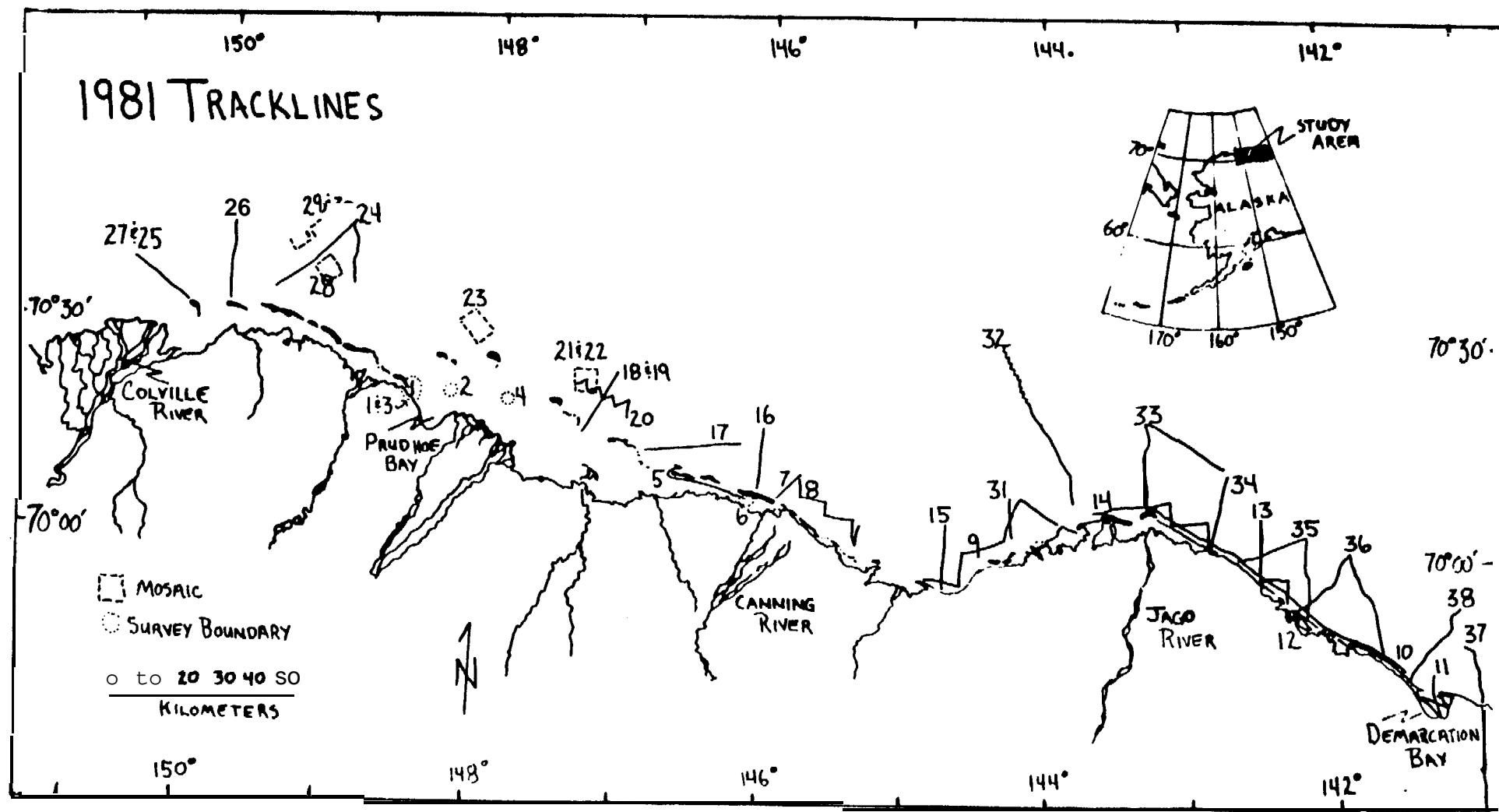


Figure 1. 1981 geophysical tracklines and site specific surveys, with line numbers listed in Table I.

Del Norte Trisponder system with a distance measuring accuracy of **+3 m**. This system provides a position accuracy of **+8 m**.

Miscellaneous studies - Three **ocean bottom** seismographs were deployed overnight at five different localities in shallow water between longitude 148° West and **the** Canadian border. The water depth ranged up to about 4 m. The purpose of this work was to monitor reported low-frequency natural seismicity in areas of decaying permafrost.

The diving investigation consisted of a roughly 1.5 **km** dive sled traverse through the area of the North **Stamukhi** Shoal side-scan **sonar** mosaic. A transponder was placed on the sea floor on each end of the traverse to facilitate rerunning of the **ship-** and diving surveys in later years. A large area around each transponder was seeded with lead birdshot for follow-up studies of sedimentation and shoal migration.

EQUIPMENT USED

Bathymetry was recorded on a Raytheon RTT 1000 dry paper recorder using either a hull-mounted 200 **kHz** transducer with **an** 8° beam width, **or** a 200 **kHz** transducer with a 4° beam width (narrow beam). All records were corrected for draft of vessel or tow depth. A 7 **kHz** transducer was used in conjunction with the RTT 1000, recording subbottom reflectors up to 10 m below the sea floor. Deeper penetration high-resolution seismic data were recorded on an EPC model 1400 recorder using 1/4 second sweep and firing rate with a 300 Joule **EG&G** Model 234 **Uniboom** as a sound source. The signal was filtered to approximately 600-1600 **Hz**.

Side-scan sonar records were taken using **a** Model 259-3 **EG&G** wet paper system and a Model 272 sonar fish with a 105 **kHz** 1/10 second pulse at a 20° beam angle depression. Records were also taken on a Model **SM** 960 **EG&G** digital system. The digital data for the mosaics were recorded on magnetic tape on a Kennedy Model 9000 magnetic tape recorder. The Model 272 sonar fish was used for both systems--the digital and the wet paper recorders.

OBS data were recorded **on** a 3-receiver system designed and built by Polar Research Laboratories of Santa Barbara, California. The three units were deployed in triangular arrays at each of 5 sites, with an internal spacing of about 100 m.

DATA ACQUIRED

Geophysical data acquired (see table 1) consist of approximately 1005 **km** of **trackline bathymetry** along with 7 **kHz** **subbottom** profiles, 800 **km** of **side-scan** sonar records, and 500 **km** of **Uniboom** seismic reflection records. **The** data listed in table 1 are keyed to figure 1. The data are in the form of 29 rolls of **bathymetry**, 20 rolls of side-scan sonar, 10 **rolls** of **Uniboom** records, 5 rolls of **Simrad** fathometer records, 38 reels of recorded side-scan magnetic tape, 120 hours of OBS magnetic tape, 8 field maps, and the ship's log. The ship's log contains important information on systems in use on each line, system settings (scale, filters, etc.), navigational data used in plotting positions, severity of ice conditions and course-holding problems and unique observations or systems difficulties. Copies of all field data are available on microfilm from the **National** Geophysical and Solar Terrestrial Data Center, NOAA, Boulder, Colorado. The microfilm is a copy of the geophysical records,

Table 1 - Geophysical data*

Line No.	Description	Raytheon	Side-scan	Uniboom	Kilometers
1	West Dock	yes	--	--	22
2	Niakuk Island	1	--	--	10
3	West Dock	1	--	--	22
4	Exxon Island	2	--	--	7
5	Outside Leffingwell Lagoon	2	--	1	24
6	Flaxman Island channel	2	--	1	6
7	Outside Flaxman Island	3	1	--	9
8	West Camden Bay	3	1	--	17
9	East Camden Bay	4	2	1	56
10	East of Jago Spit	6	5	--	81
11	Demarcation Bay	7	6	2	30
12	Beaufort Lagoon	8	7	3	17
13	Outside Beaufort Lagoon	9	7	3	29
14	East of Jago Spit to Barter Island	10	8	4	43
15	Test Line 7	11	9	4	19
16	Test Line 8	12	9	5	17
17	East of Pole Island	12	--	--	26
18	Test Line 6	13	10	5	17
19	Reciprocal, Test Line 6	13	11	--	17
20	18-m bench delineation	14	--	--	28
21	Mosaic northeast of Narwhal Island	15	12	--	55
22	Continue mosaic	16	12	--	
23	18-m bench north of Reindeer Island	16	13	--	23
24	Cat Shoal	17	--	--	45
25	Test Line 1	--	13	--	10
26	Test Line 2	18	14	--	20
27	Test Line 1	18	14	--	19
28	South Stamukhi Shoal Mosaic	19	14	--	46
29	North Stamukhi Shoal Mosaic	21	16	--	
30	Rerun 1977 lines on Stamukhi Shoals	23	--	--	65
31	Camden Bay to Barter Island	23	--	6	9
32	Continental Shelf Run off Barter Is.	23	17	6	48
33	Seaward leg offshore east of Barter Island (+ 14 km run over)	25	18	7	20
34	Shoreward leg east of Barter Island	26	18	7	19
35	Dogleg offshore & back into Pogok Bay	26	19	8	41
36	Offshore and back outside Beaufort Lagoon	27	19	8	52
37	Line at U.S./Canadian Border	29	20	9	19
38	Offshore Demarcation Bay	29	20	10	24

*Numbers in the Raytheon, side-scan and uniboom columns represent beginning roll numbers and signify data gathered on that line by that system. No number means the system was off.

ship's log and computer print-out of digitized way points. The printout of these way points **would** allow for reproduction of tracklines at any scale, and correlation to geophysical records through time points. Originals are archived at the U.S. Geological Survey, Deer Creek Facility, **3475 Deer Creek Road**, Palo Alto, California 94304.

Surface samples collected are listed *in table 2*, along with water depth, longitude, and latitude, and shipboard sample descriptions and observations. The locations are shown in figure 2. Almost all samples were obtained with a grab, and cuts from most were given to Dr. Bill **Briggs** for studies of Ostracodes. The bulk of the material is being kept at our facility in Palo Alto, California.

DATA ANALYSIS

In our analysis of the geophysical reconnaissance data obtained between the Canning River and the **Canadian** border the focus has been on ice gouging. For the analysis we have basically used the shore-normal transects and eliminated shallow-water, shore-parallel lines (Fig. 3). The short time available for **analysis** required reduction of the number of parameters extracted from monographs and fathograms, compared to the very thorough analysis completed for the region west of the Canning River (**Rearic** et al. , 1981) . A copy of the completed data sheets used in this study is presented here as an Appendix. As in previous **work**, the **tracklines** are broken into 1-km-long segments, as listed in the first column of the data sheets. The parameters we considered most significant for this study are the following:

1. Water depth - to find relationships to severity of gouging.
2. Gouge depth - maximum gouge incision depth per km segment.
3. Ridge height - to allow calculation of total relief from gouging.
4. Gouge width - maximum per km segment.
5. Gouge density - the number of gouges actually counted is to the left of the normalized count listed in this column and separated by a slash.
6. Gouge orientation - dominant trend with respect to trackline is to the left of the true north orientations and separated by a slash.
7. Sediment cohesion - an attempt to judge from geophysical records whether the bottom is composed either of sand and **coarser** non-cohesive material, or of fine and cohesive material, as reflected in the shape and character, of the gouges.

We also measured the depth below sea level of the first continuous **subbottom** reflector seen on the 7kHz records ("Reflector A"). The main purpose of extracting this data was an attempt to relate ice gouges to the geology of the shelf surface. Subtracting water depth from "Reflector A" gives what we consider the maximum possible thickness of Holocene marine sediments blanketing the shelf. Given the fact that ice gouging has repeatedly disrupted the sediments since the last transgression, the Holocene marine sediments should not contain continuous internal reflectors in seismic records, an assumption that has strong support from detailed studies done in the **Prudhoe** Bay region. But until more detailed work allows us to correlate through the entire region of this **reconnaissance** survey, we cannot put much emphasis on this data.

TABLE 2

Station Number	1981 Sample Descriptions					
	Latitude	Longitude	Water Depth (m)	Type Sample	Reference Location	Description
4	70.387°	148.515°	2	achunl	W. Dock	Core length 37.5 cm. Very thin souo on top overlying mud with-gravelly mud at-base. "
5	70.1050	145.324°	15.5	Grab	Line 8	On seaward flank of shoal. Sand
6	70.1040	145.3260	12.5	Grab	Line 8	On seaward flank of shoal. Clean sand
7	70.1030	145.328°	9.5	Grab	Line 8	On crest of shoal. Coarse sand
8	70.1020	145.3300	13	Grab	Line 8	Inside shoal. Coarse sand and pea gravel 1-2 cm) over grey mud.
9	70.1010	145.3330	13	Grab	Line 8	Inside shoal. Sandy mud. Few pebbles
10	70.0200	145.3150	n beach		Landen Bay	Outcrop of stiff silty clay (?)
11	69.6750	141.3190	5	Grab	Demarcation Bay	Sandy mud with bivalves.
12	69.6560	141.2810	4	Grab	Demarcation Bay	Organic mud, silt and clay with trace of sand .
13	69.6550	141.3560	4	Grab	Demarcation Bay	Organic mud w/worm tubes.
14	69.8590	142.1630	2.2	Grab	eaufort Lagoon	Sandy organic-rich mud. Peaty material - brown. to black
15	69.8900	142.2530	3	Grab	eaufort Lagoon	Muddy organic sand
16	69.9090	142.3150	2.5	Grab	eaufort Lagoon	Muddy organic sand
17	70.1270	142.5000	35	Grab	Offshore Pokok Bay	Muddy sand. Soft!
18	70.0560	142.4880	23.5	Grab	Offshore Pokok Bay	Sandy mud
19	70.0310	142.5360	18.5	Grab	Offshore Pokok Bay	"After 3 lowerings muddy gravel. Gravel w/benthic growth Stiff, silty clay below?
20	70.0170	142.5220	16	Grab	Offshore Pokok Bay	Fairly clean sand overlain by 1-2 cm of muddy sediments.
21	69.9890	142.5180	7	Grab	Offshore Pokok Bay	Clean fine sand

TABLE 2

1981 Sample Descriptions

Station Number	Latitude	Longitude	Water Depth (m)	Type Sample	Reference Location	Description
22	70.633	148.1600	...	Ice	N. of Reindeer	Stamukhi ice
23	70.6330	148.1690		Ice	N. of Reindeer	Gravelly mud on only one surface of blocky ice floe.
24	70.620 ^c	148.1270	18	Grab	18-m bench/Reindeer	Crest of ridge. Muddy gravel, overconsolidated?
25	70.620 ^c	148.1460	18	Grab	18-m bench/Reindeer	Samples 24, 25, 26 at top of break in slope on 18-m bench
26	70.620 ^c	148.1670	18	Grab	18-m bench/Reindeer	all muddy gravel of various consistencies, from soupy on the west to stiff on the east.
27	70.498 ^c	143.203 ^c	52	Grab	Line 32	Gravel, up to 3 cm diameter w/bryozoans and other small growth in big gouge terrain with rounded relief. Between pebbles apparently is a trace of trapped transient mud.
28	70.357 ^c	143.2920	40		Offshore Barter Is	Medium firm grey mud w/ a few scattered very small pebbles.
29	70.2300	142.7470	40	Grab	Offshore Pokok Bay	Firm mud w/ a 5-cm layer of soft mud on top. No shells or pebbles.
30	69.8730	141.7170	23	Grab	Line 36	Pebble rich, sandy mud, soft. Pebbles up to 5 cm w/coral growth, bryozoans.
31	59.8820	141.1470	34	Grab	Line 38	Soft mud, perhaps even transient layer separated by thin black line from finer mud below. No pebbles, probably no sand.
32	59.8850	41.2420	32	Grab	Line 38	Slightly silty clay, increasing very gradually from soupy on surface to slightly firmer below. Several small shells, no pebbles.
33	59.8160	41.2590	30	Grab	Line 38	Silty clay, grey as sample 32 w/gradual increase in strength downward, no sand, small brittle star.
34	59.7860	41.3700	23	Grab	Line 38	Slightly pebbly, sandy mud. Soft at surface (5 cm) and firmer at bottom (15 cm).
35	69.7540	41.4440	6.5	Grab	Line 38	Pebbly, slightly muddy sand. One large pebble (6 cm), subrounded, with much growth, including bryozoans, coral, etc.
36	69.7390	41.4640	2.5	Grab	Line 38	Clean pebbly sand, one clast 6 cm. No growth, no mud.
37	69.7190	41.4790	7.5	Grdb	Line 38	After 3 borings: muddy gravel, clast to 10 cm, no growth.

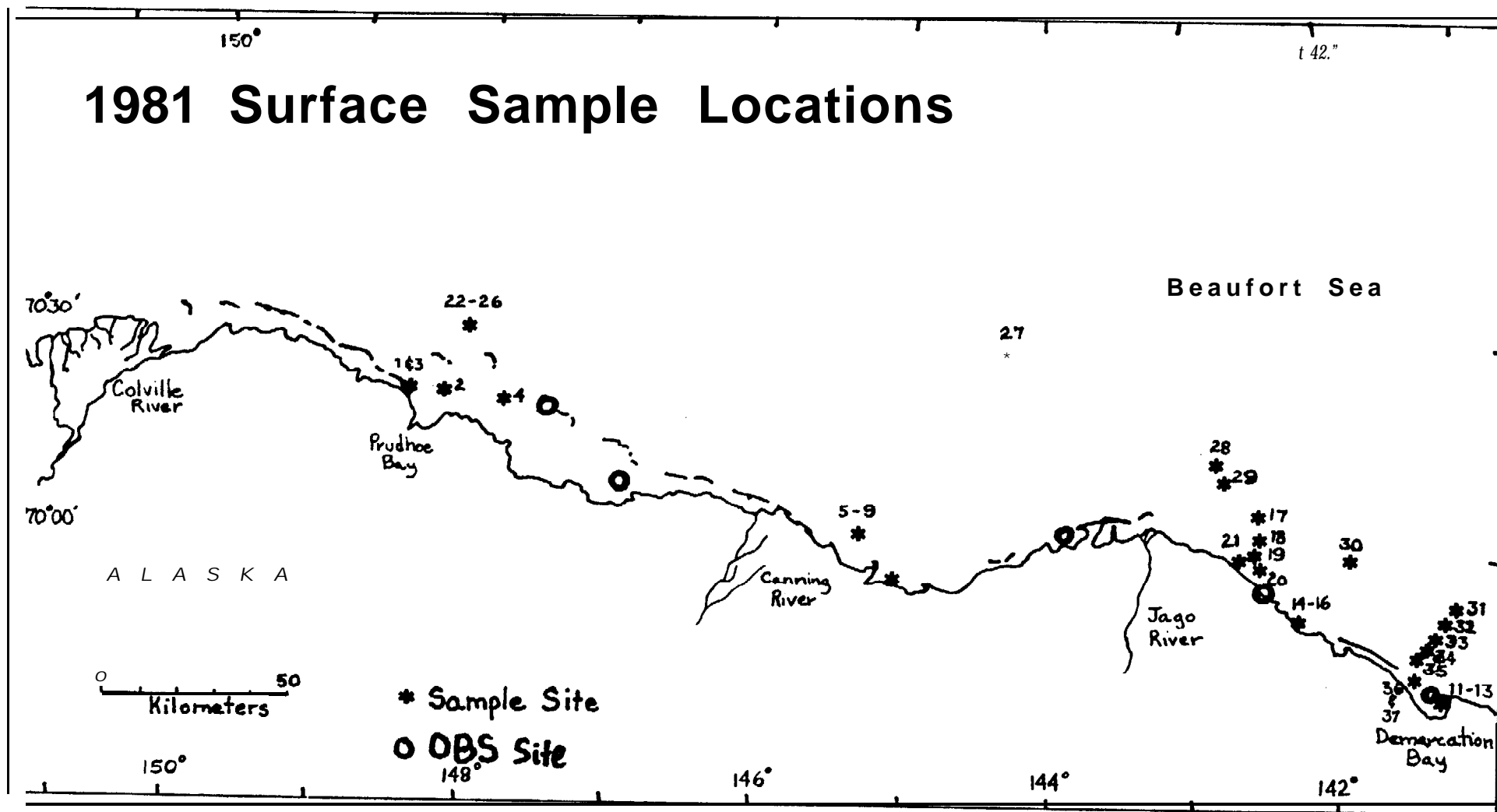


Figure 2. 1981 station locations for grab samples and Ocean Bottom Seismographs (OBS).

A computer was used for plotting certain gouge parameters on maps, for simple statistical analyses, and for preparing scatter plots of gouge parameters.

RESULTS

Bathymetry - The bathymetry shown in figure 3 *is* from Greenberg, et al. (1980), and we generally found no major disagreements with the water depths recorded along our **tracklines**. However, the trackline off the Canadian border should have crossed a broad shoal suggested by published data, but we found no indications for this feature. Previous **work** has shown the important role played by shoals in ice dynamics and in controlling ice zonation (Rearic and Barnes, 1980; Reimnitz et al., 1978), and we therefore indicate the major topographic highs crossed by our survey lines, along with the height above the surrounding sea floor. We assume that these features are oriented generally shore parallel as suggested in figure 3. Only the shoal off the Canning River was surveyed by a zigzag trackline pattern and is well defined.

Ice Gouging - The pattern of dominant ice gouge alignment parallel to regional **isobaths** as mapped west of the Canning River (Barnes et al., 1981) continues eastward to the Canadian border (Fig. 4). The Barter Island region, forming a major promontory jutting out into the pack-ice drift of the clockwise rotating Beaufort Gyre, separates **two** regions with distinctly different **isobath** trends and ice-gouge trends. In figure 5 we plotted **water** depth against **dominant** gouge orientation. A clear break is shown at 18-20 m water depth, with considerable orientation scatter shoreward, and parallel alignment seaward. The mean gouge orientation of 103°T in the study area is heavily weighted by trend determinations corresponding to the NW-SE trending isobaths east of Barter Island. By comparison, the mean gouge orientation west of the Canning River is 90°T.

Ice gouge density values (adjusted gouge counts per km of **trackline**) have been contoured in figure 6. A very well defined zone with over 150 gouges per km of **trackline** lies in water 18-36 m deep. This zone has been defined by Reimnitz et al. (1978) as the stamukhi zone. The **scattergram** (Fig. 7) shows a clear trend of increasing gouge densities from the shore to the stamukhi zone, and decreasing gouge densities from there to 58 m water depth. The greatest depth at which a gouge was seen was at 58 m on **line** 32, which extends to the edge of the shelf. The mean gouge density in the survey area is 108, compared to a value of 63 for the region west of the Canning River. We believe that these higher gouge counts are explained largely by the fact that mean water depth for the areas surveyed here is 25 m, whereas west of the Canning River the mean depth is 17 m.

The maximum gouge incision depths have been contoured in figure 8. Again the 18 m **isobath** is a dividing line between maximum incision depths of less than 1 m inshore and greater than 1 m offshore, as also shown on the scattergram in figure 9. But the maximum incision depths and the maximum gouge widths (Fig. 10) continue to increase seaward and do not begin to decrease until the very outer ice-gouge limit observed on lines 32 and 33. The mean for all maximum incision depths in the study area is .8 m, compared to .5 m for the western region. The mean of the maximum incision widths is 10 m, versus 8 m for the western region. Again the larger gouge size can be explained in part by the greater average water depth in the present study area.

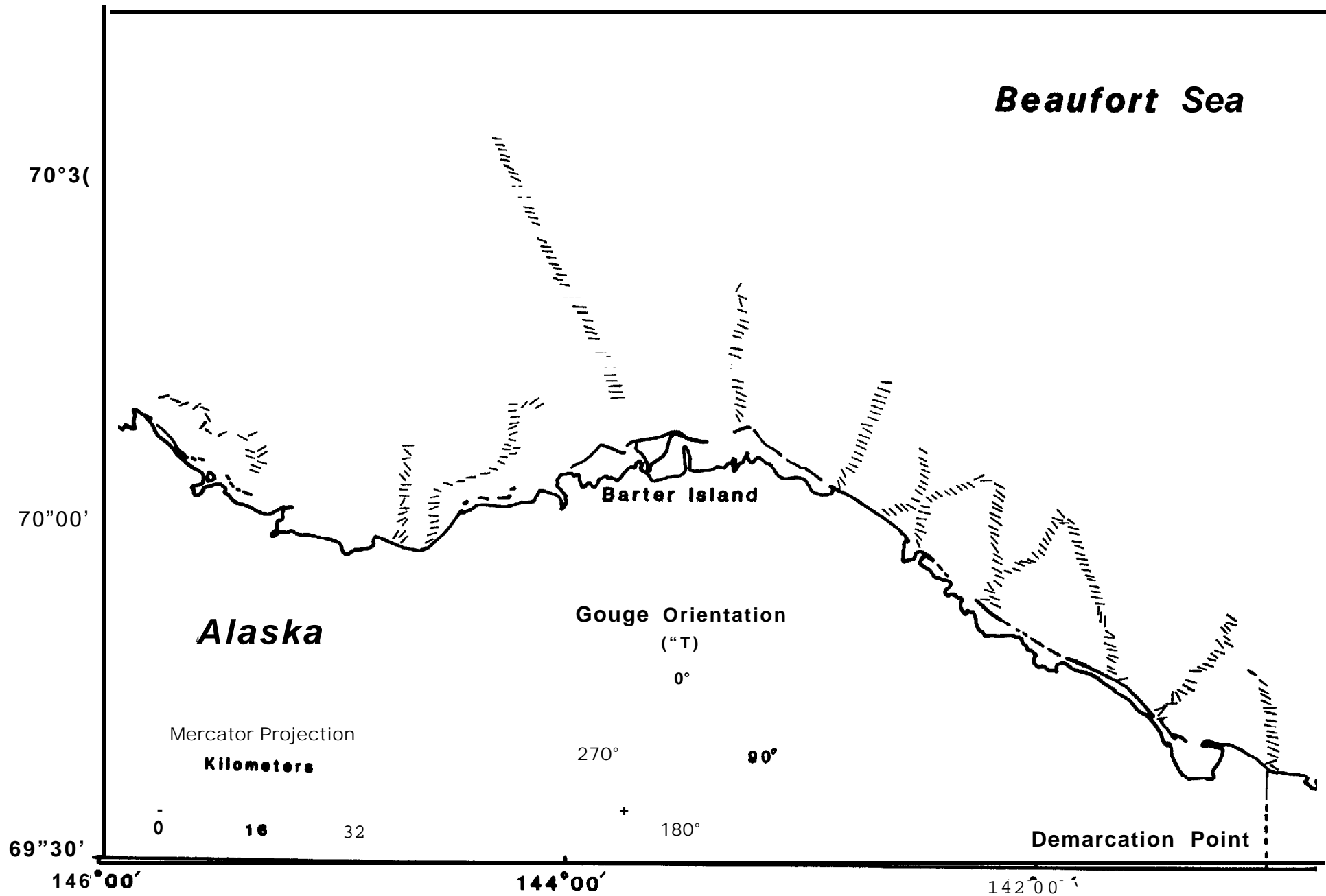


Figure 4.- Gouge orientations in the Barter Island area. Each line represents the dominant gouge orientation measured over 1 km of **trackline**.

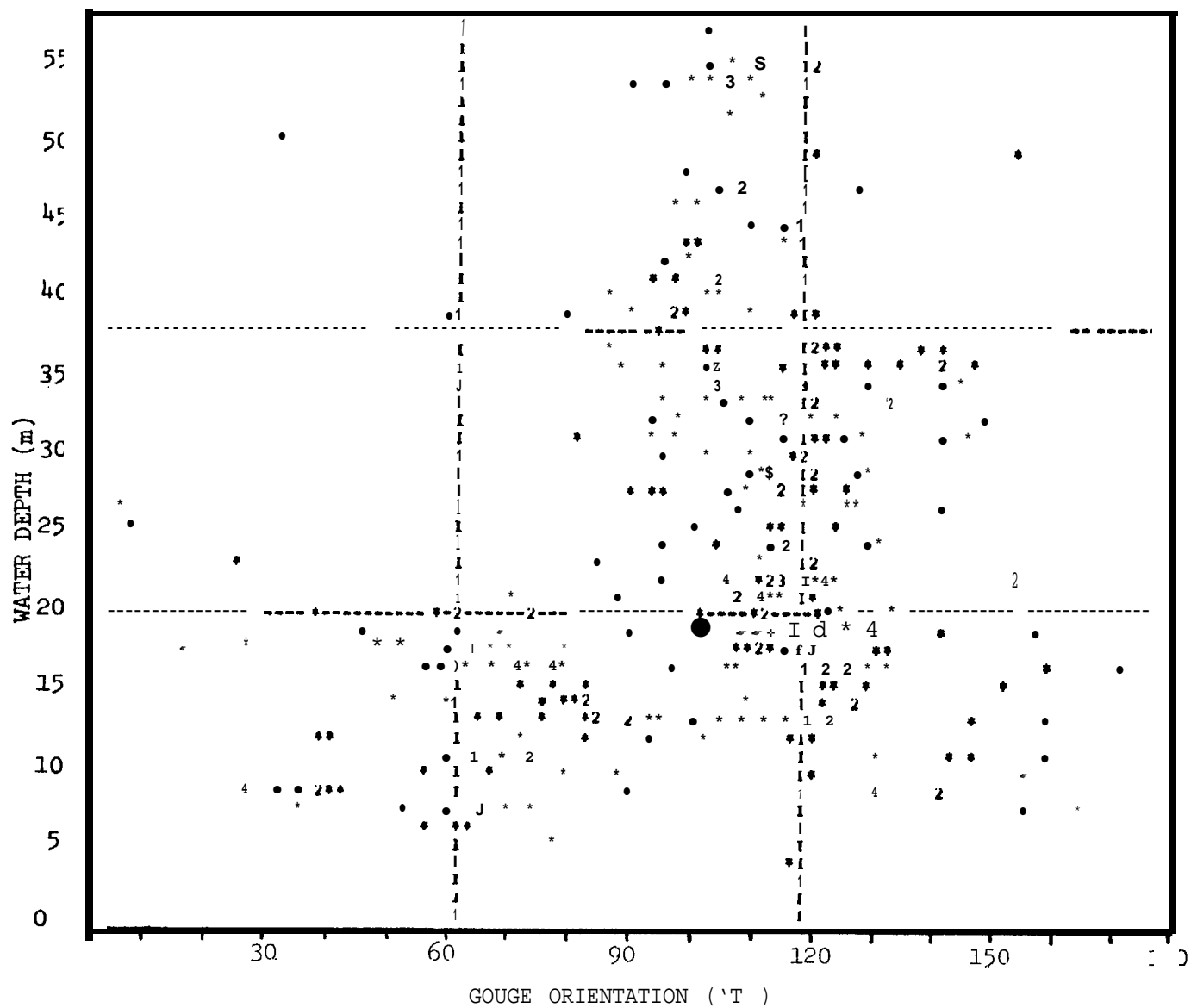


Figure 5.- Scattergram of gouge orientation versus water depth, showing wide scatter at water depths shallower than 18 meters.

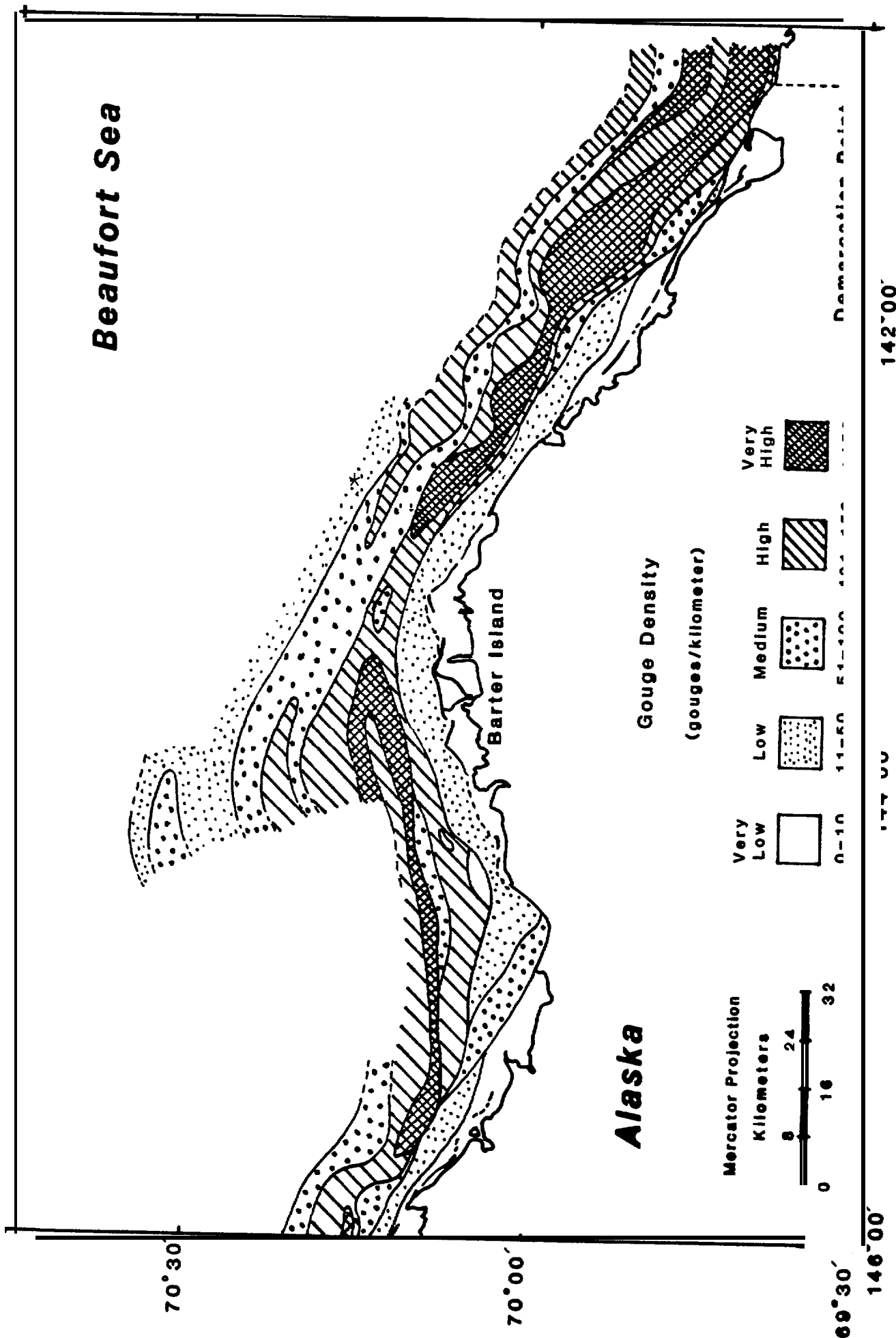


Figure 6.- Contours of ice gouge density values from Camden Bay to the Canadian border.

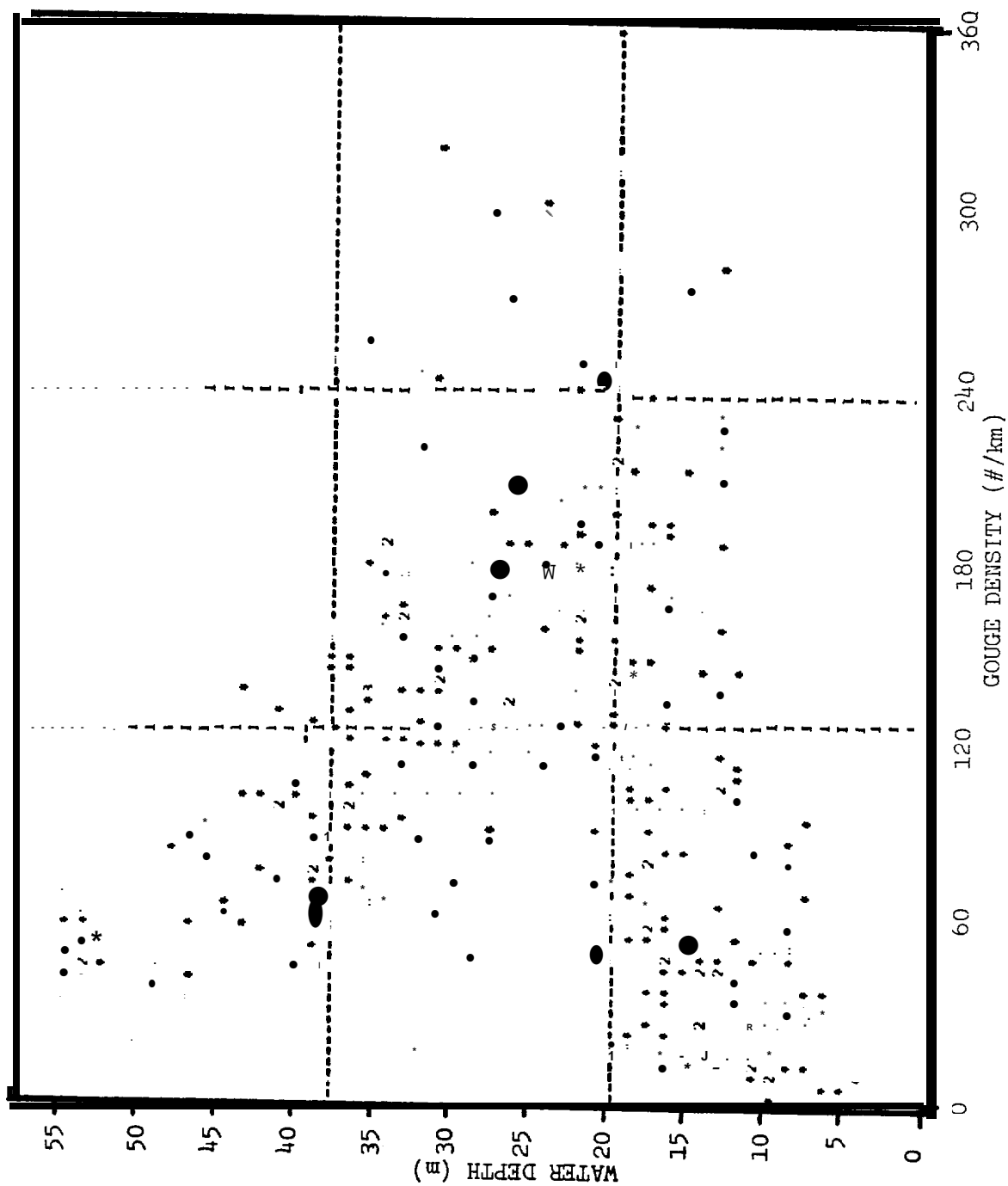


Figure 7.- Scattergram of gouge density versus water depth.

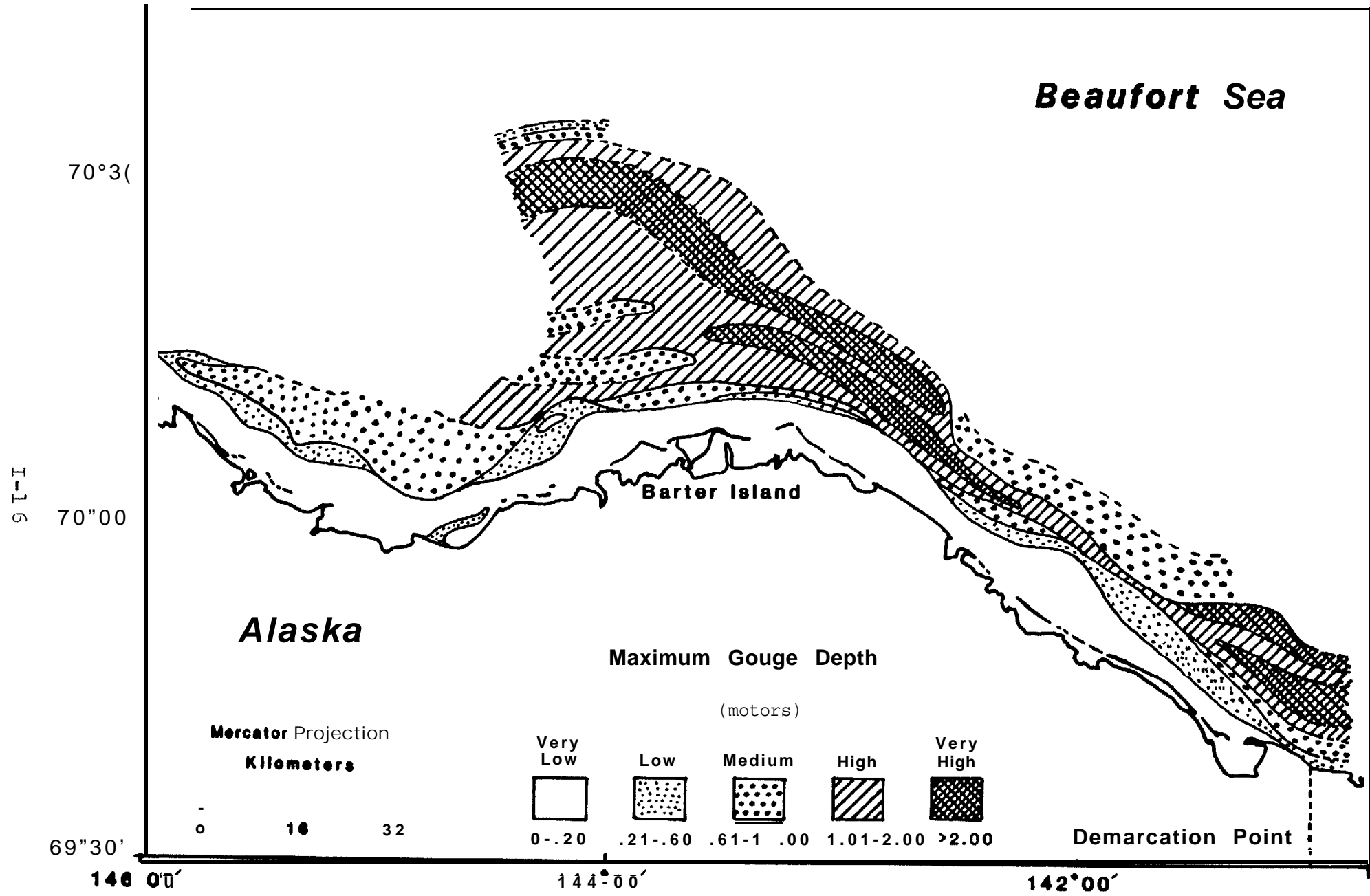


Figure 8.- Contour map of ice gouge maximum incision depth for the area from Camden Bay to the Canadian Border.

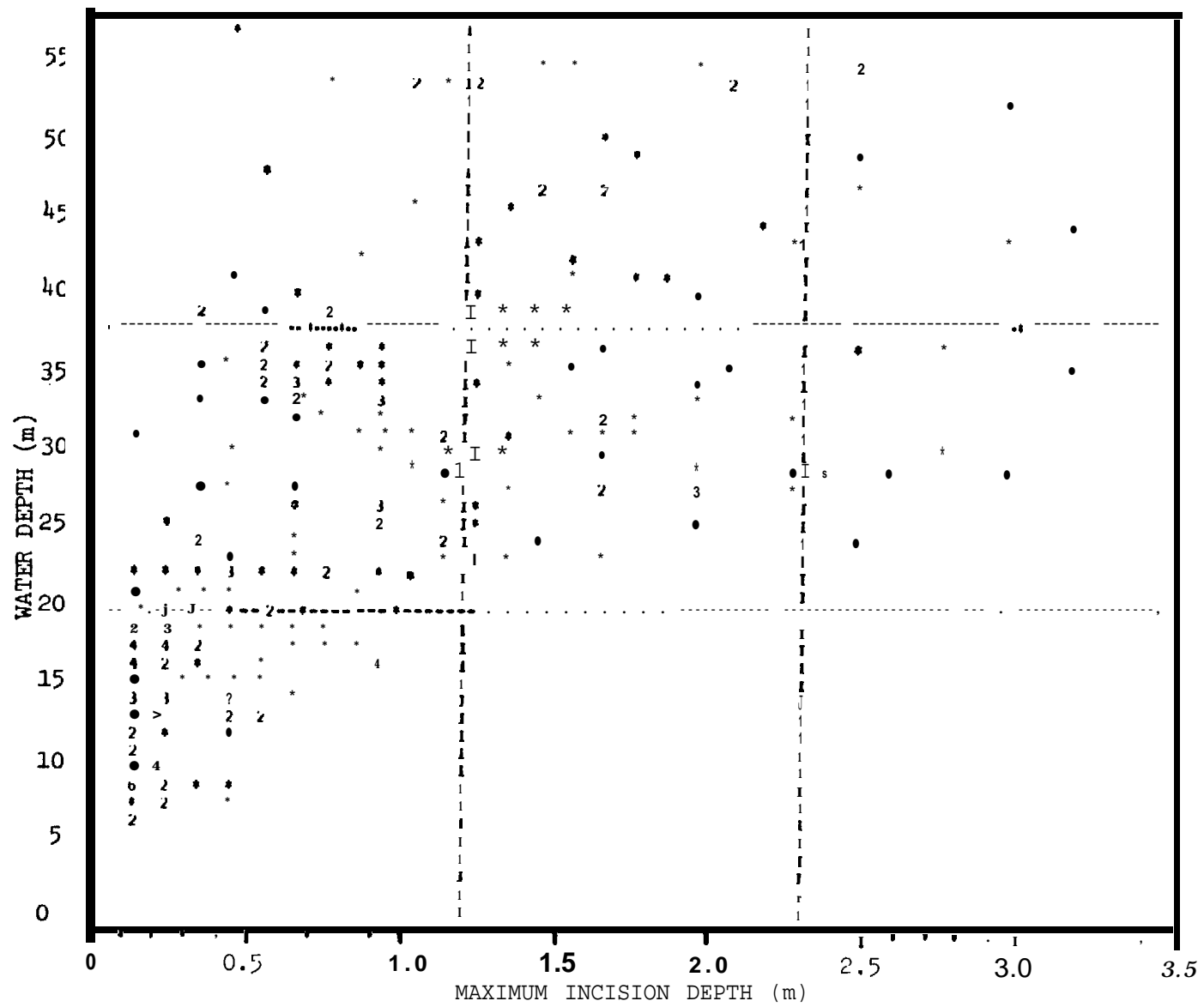


Figure 9.- Scattergram of ice gouge maximum incision depth versus water depth.

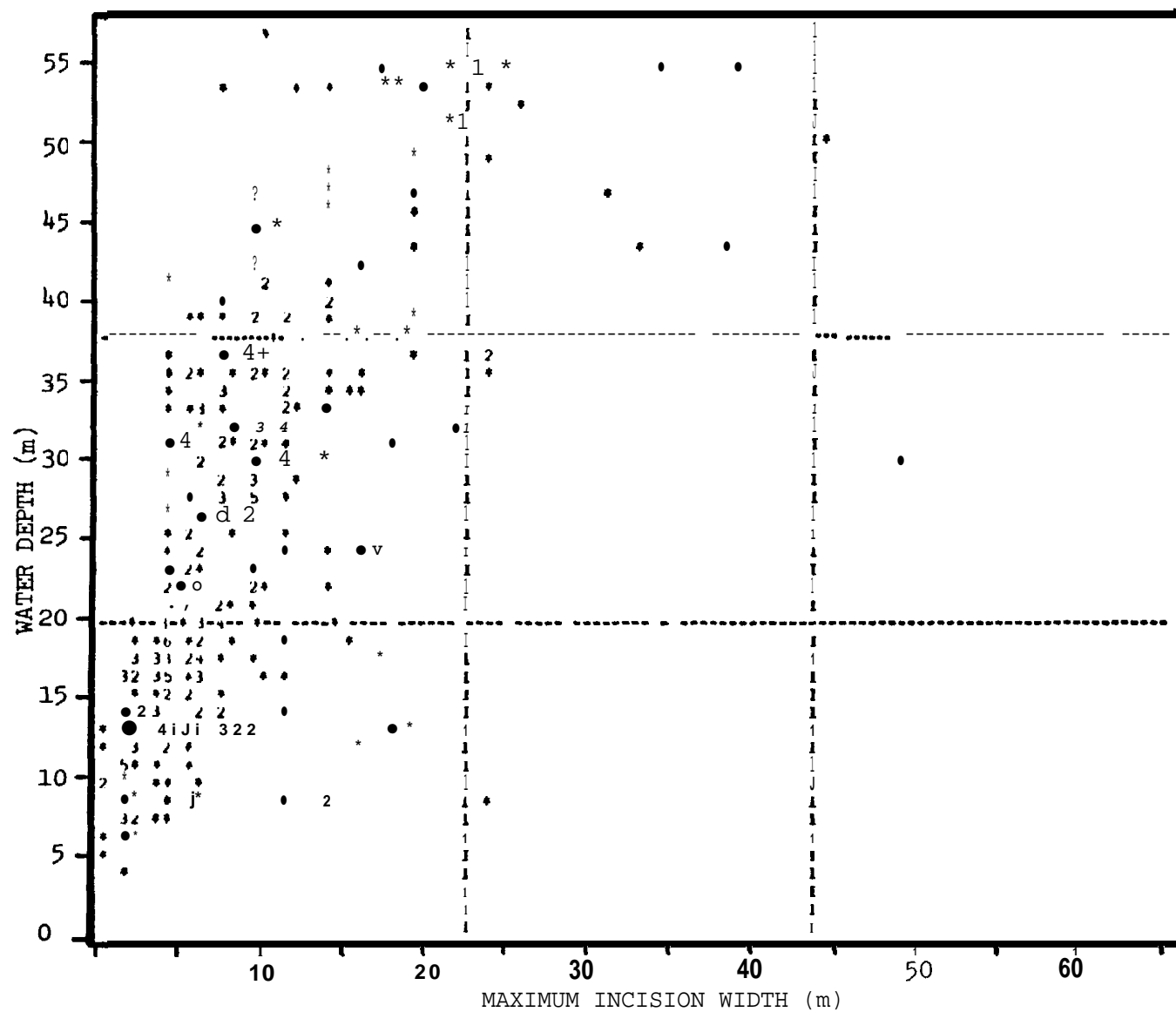


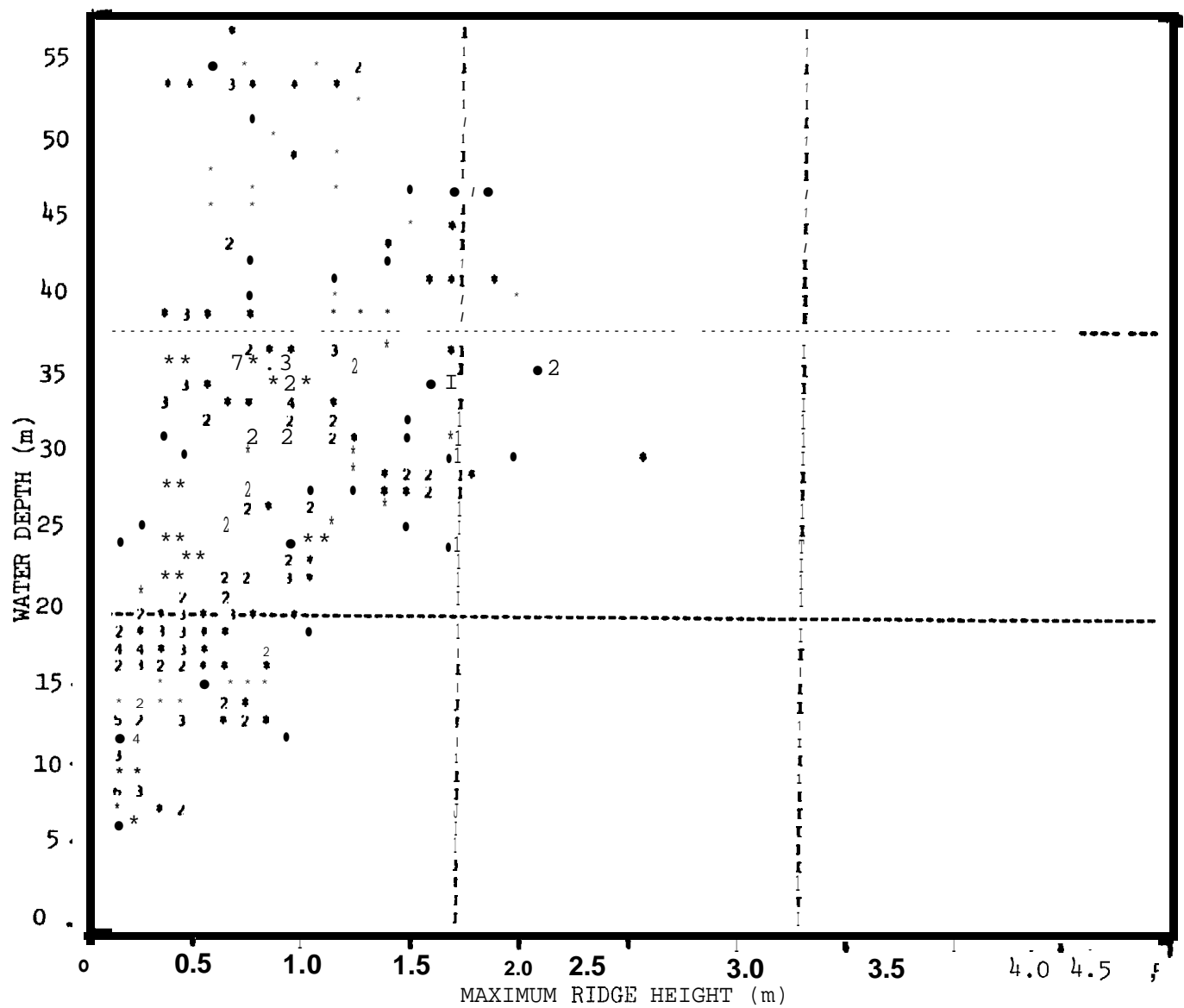
Figure 10. - Scattergram of ice gouge maximum incision width versus water depth.

Figure 11 is a scattergram of ridge height versus water depth. This shows that shoreward of the **18 m** isobath ridges are no higher than 1 m. Ridges are highest between the 25-m and 45-m **isobaths**, and decrease from there seaward. This is contrary to the continuous increase in gouge depth and width measurements with increasing water depth. Total ice gouge relief (incision depth plus ridge height) was plotted against water depth in figure 12 and shows an increase offshore with a slight drop near the outer limit of ice gouging. Barnes et al. (1980), based on the highest ridges and greatest incision depths seen in the western area, speculated that total **relief** could reach 8 m in a single gouge. In the present study the greatest **value** for total relief seen in a single gouge was 8 m and found in water 38 m deep.

Figures 13 and 14 are **scattergrams** of gouge density plotted against maximum incision depth and maximum incision width respectively. Both scattergrams show that with increasing gouge density there is a corresponding decrease in gouge size. This inverse relationship can be explained by the fact that large gouges take up more space in each counting interval than smaller gouges and correspondingly fewer large gouges can be fit into such an interval. Many small gouges may also be **reworked** by formation of one large gouge. Figure 15 shows a plot of gouge orientation versus gouge density. The difference in scatter between figures 5 and 15 demonstrates that orientation is related to water depth but not to gouge density.

Seismic reflection studies - The central portion of the study area is interpreted by Grantz and Dinter (1980) as being tectonically and seismically active and undergoing uplift during the Holocene. The geology here is more favorable for seismic profiling than in most of the regions west of the Canning River, where the data is very difficult to decipher. Figure 17 is a sample **Uniboom** record (for location see figure 16) on which the most prominent sets of reflectors have been enhanced with inked lines. A major angular unconformity lies at a depth of 10-12 msec below the sea floor. Only 3 msec below the sea floor a discontinuous faint reflector can be traced. (Assuming a sound velocity of 2,000 msec in sediment, 1 msec is 1 m on this record.) Figure **18** is a sample **Uniboom** record with the angular unconformity at the sea floor possibly overlain by an extremely thin veneer of soft sediment that cannot be traced on this record. The hyperbolic patterns within the upper 10 msec of the record are a result of the ice gouge relief on the shelf surface. We do not know whether these gouges are cut into the old dipping strata truncated by the sea floor, or whether scouring by ice has resulted in a thin residual deposit in which the gouges are formed.

Very thin surface sediment layers are best resolved on the 7 kHz record. Examples of these records are shown in figure 19 (A and B). In figure **19B** the strong dark band 1 m below the sea floor, and precisely conforming to the ice gouge relief, is the 7 kHz trace of the sea floor. The faint reflector at about 58 m below sea level is a real **subbottom** reflector. All such shallow reflectors were traced from the 7 kHz records at a very shortened horizontal scale, giving a high vertical exaggeration, and are presented as figures 20 through 23. **Tracklines** and figures are arranged **in** order from the Canning River to the Canadian border and all lines are oriented with the shoreward (S-SW) end on the left side, except tie line 33-34, which parallels the slope.



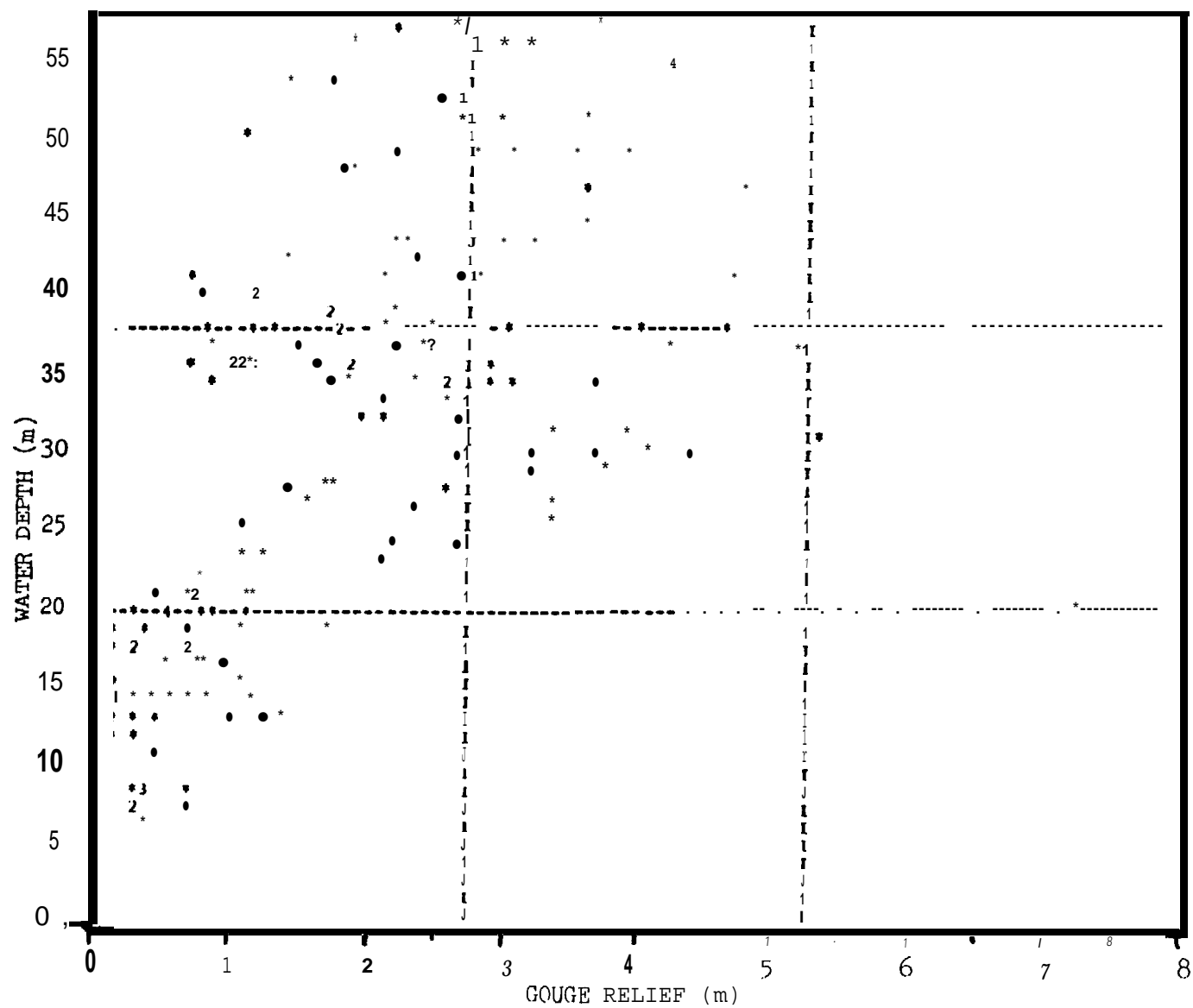


Figure 12. - Scattergram of gouge relief (ridge height plus incision depth) versus water depth.

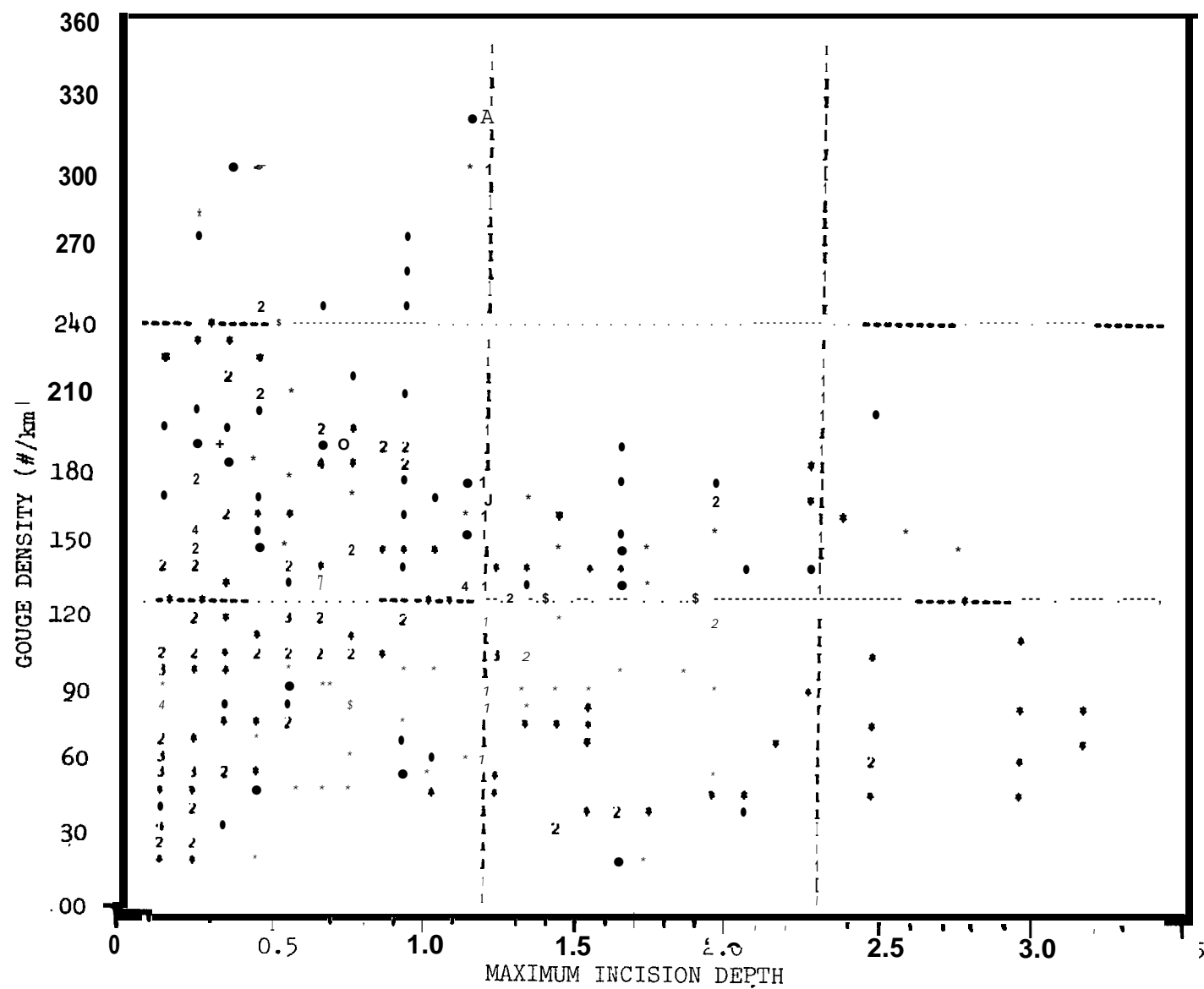


Figure 13. - Scattergram of maximum gouge incision depth versus gouge density.

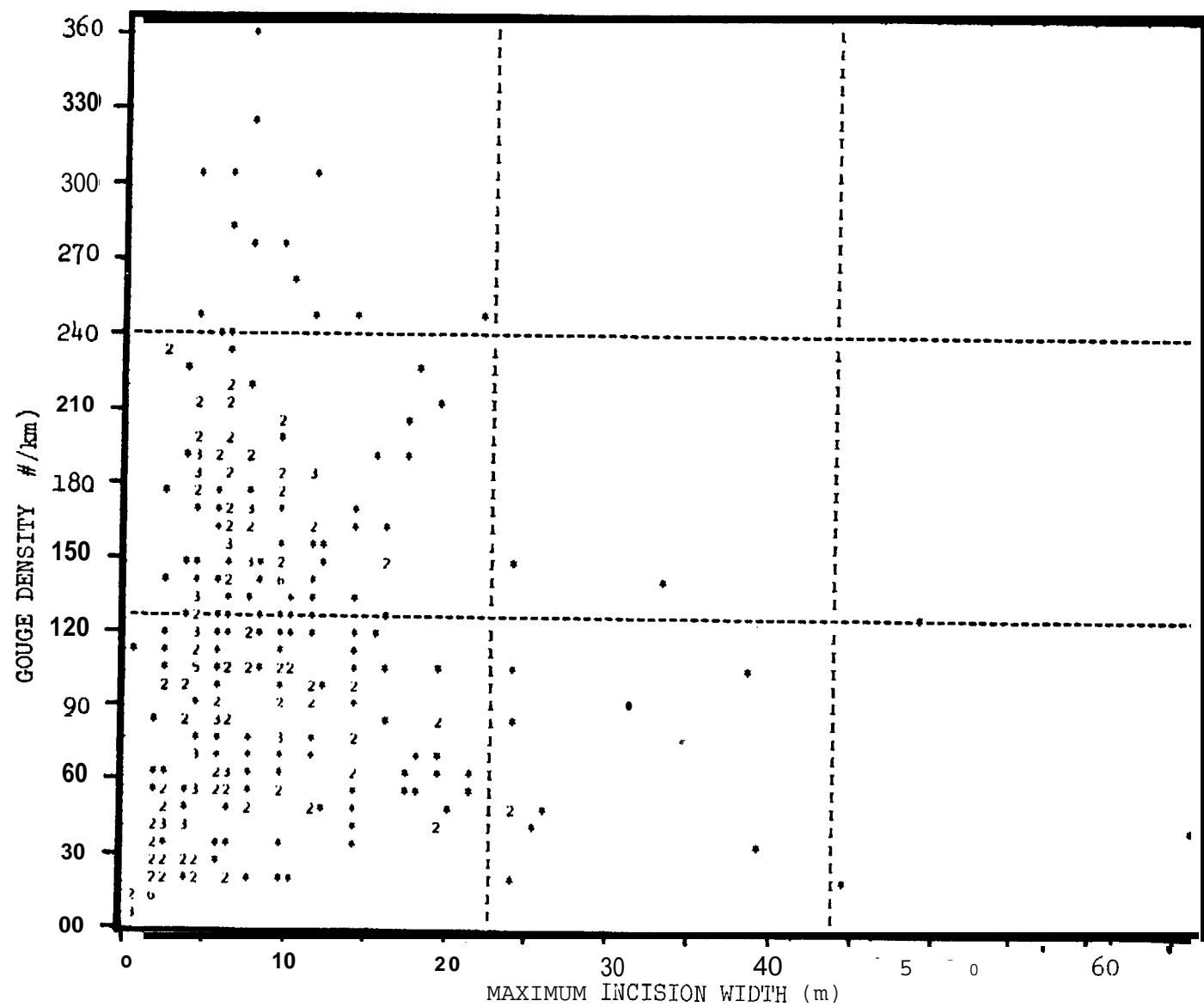


Figure 14. - Scattergram of maximum gouge incision width versus gouge density.

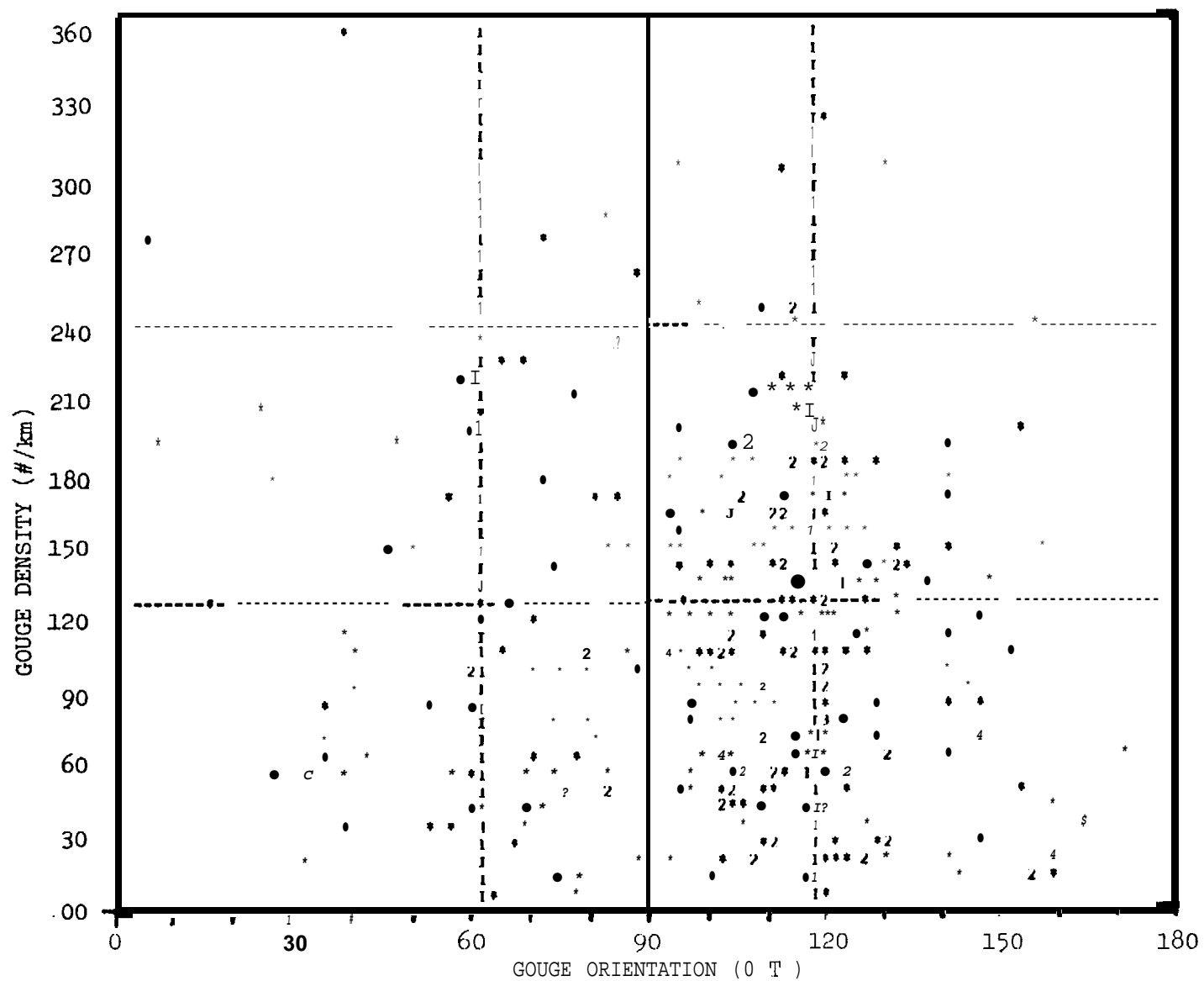


Figure 15. - Scattergram of gouge orientation versus gouge density.

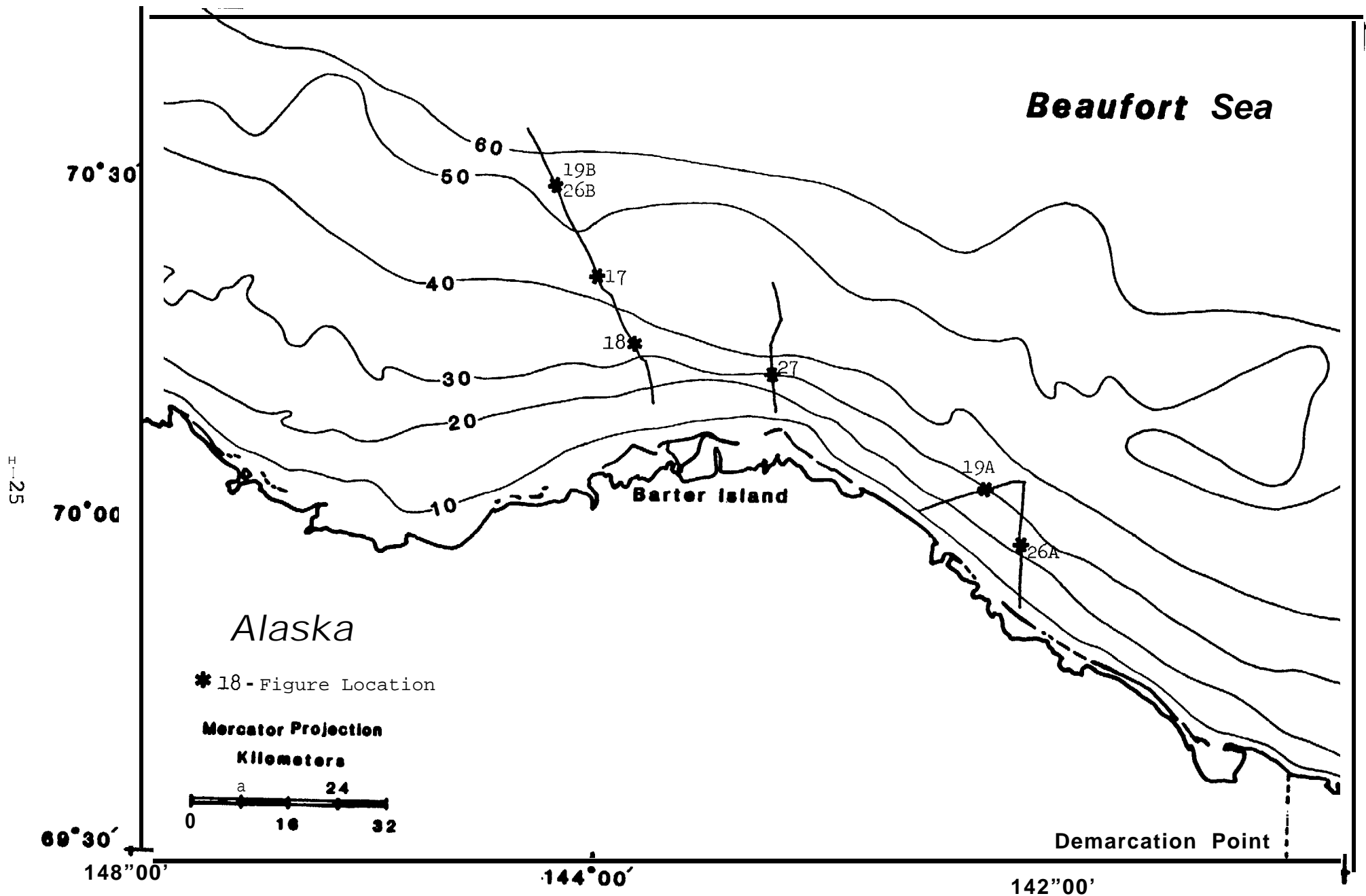


Figure 16.-Locations for fathograms and monographs shown in Figures 17,18,19 A and B, 26 A and B, and 27. The tracklines from which these examples stem are also shown.

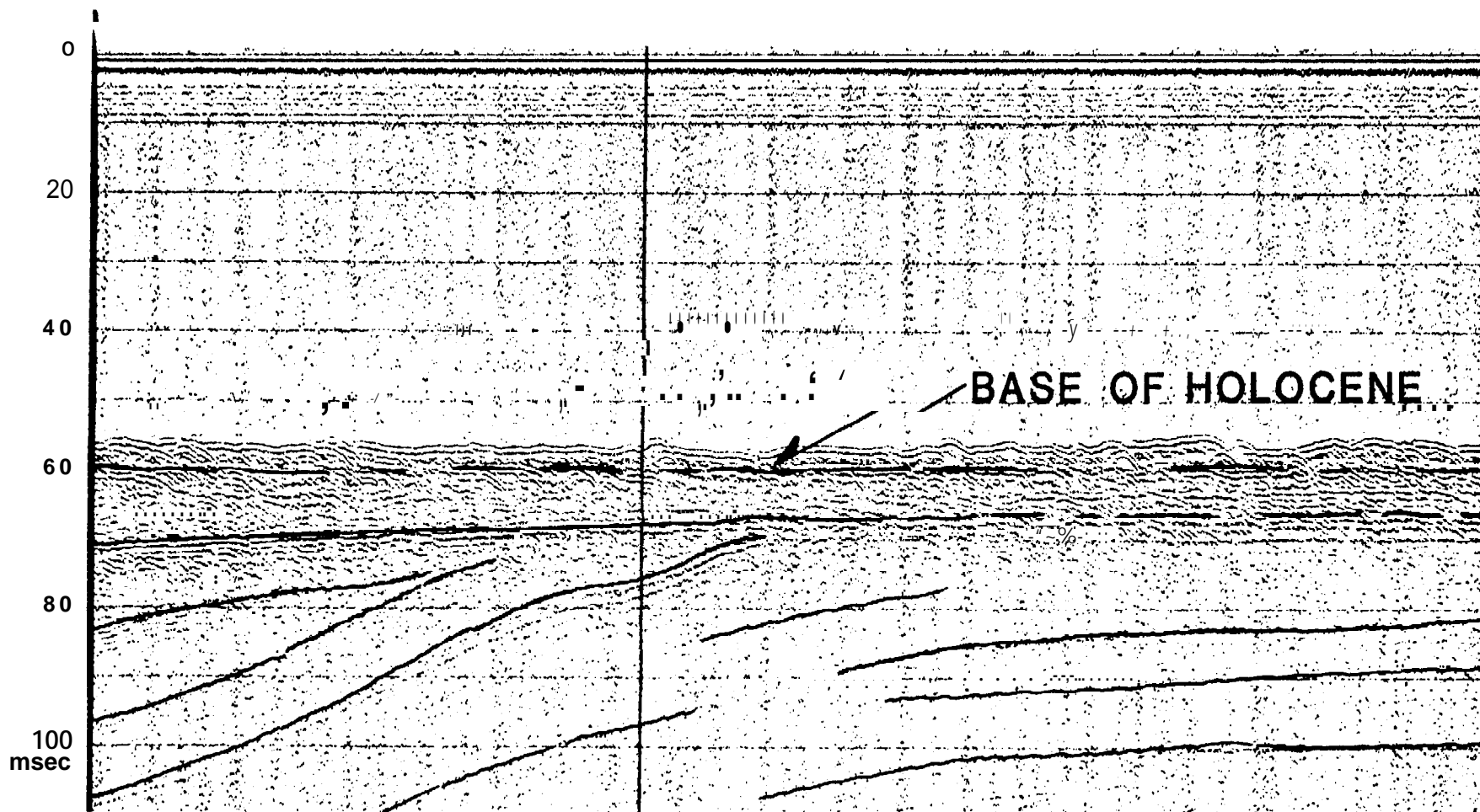


Figure 17. - High resolution **siesmic** record showing an angular unconformity below the base of the Holocene.
Some of the reflectors have been highlighted for clarity.

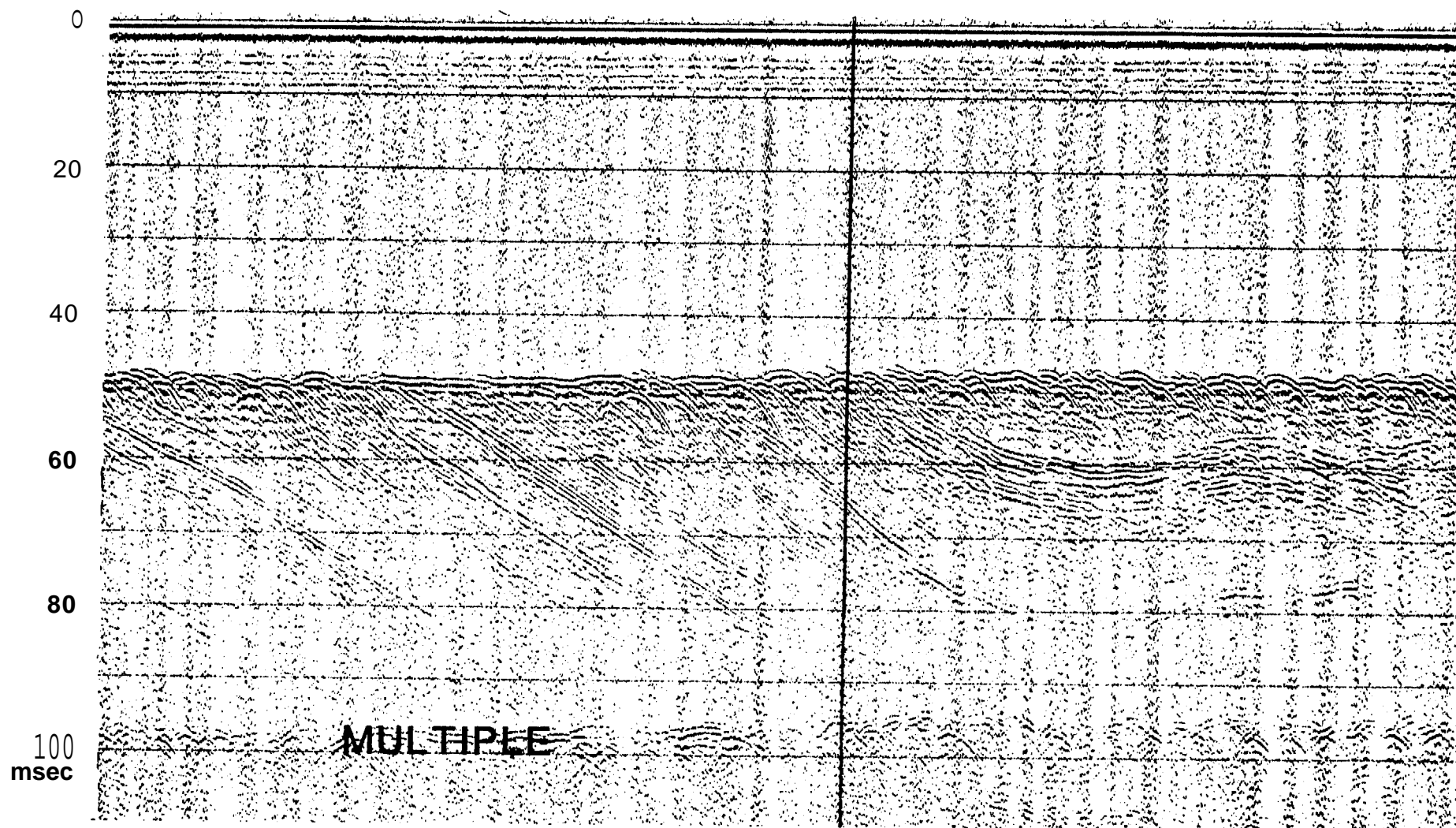


Figure 18. - High resolution seismic record showing dipping sediments truncated by the seafloor. The hyperbolas seen in the first 10 msec below the seafloor, and the rough seafloor relief, are the result of ice gouging.

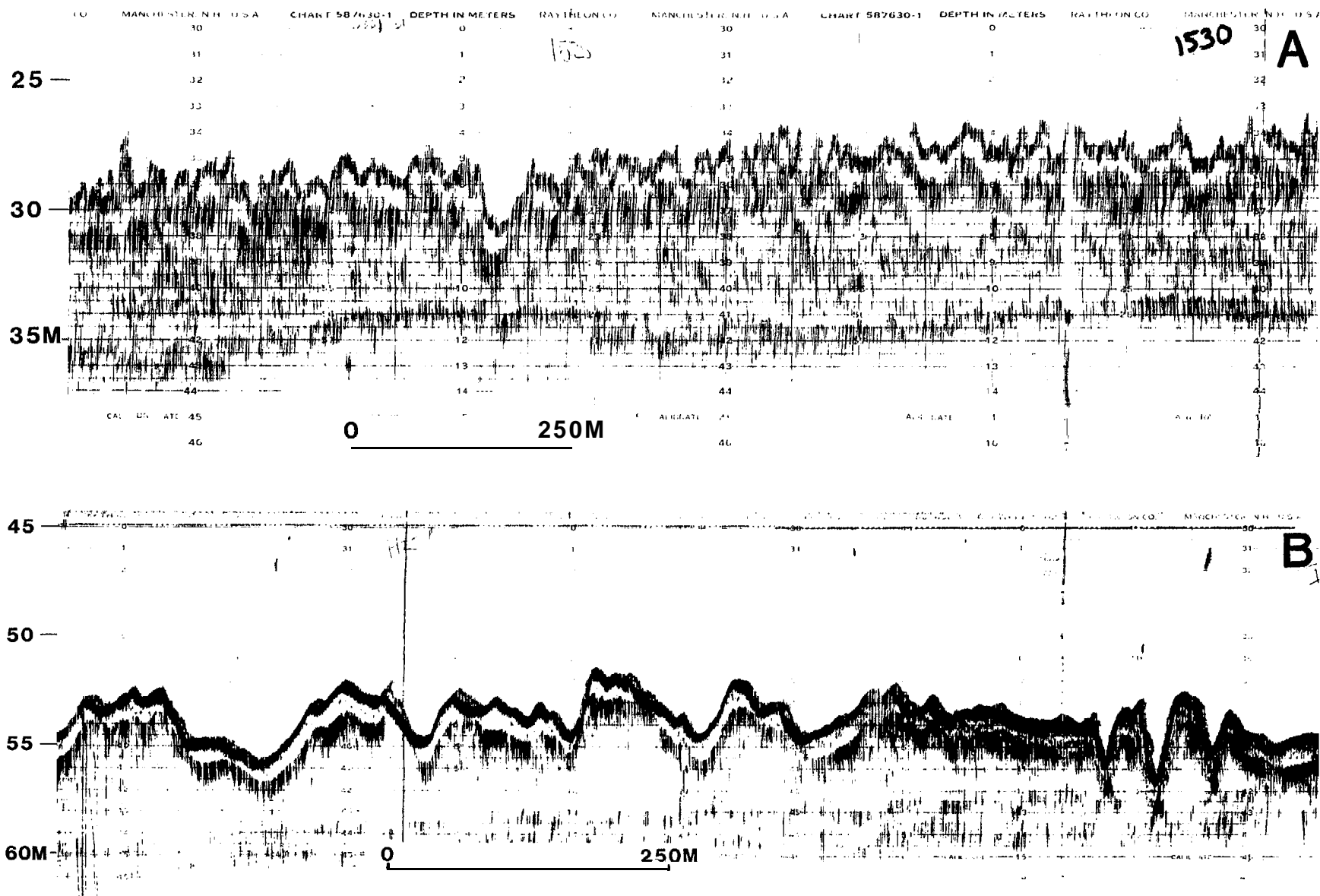


Figure 19. - Fathograms showing the extremes in bottom character: (A) rough relief and (B) smooth relief, which we interpret as cohesive and non-cohesive sediments, respectively. An undulating sub-bottom reflector is visible at approximately 35m in fathogram A, and a faint, left-dipping sub-bottom reflector is visible at approximately 60m in fathogram B. Fathogram B covers the same section of trackline as sonograph B in Fig.26. (See Figures 16 and 26.)

Surface Sediments: Cohesive — Non-cohesive - - - -

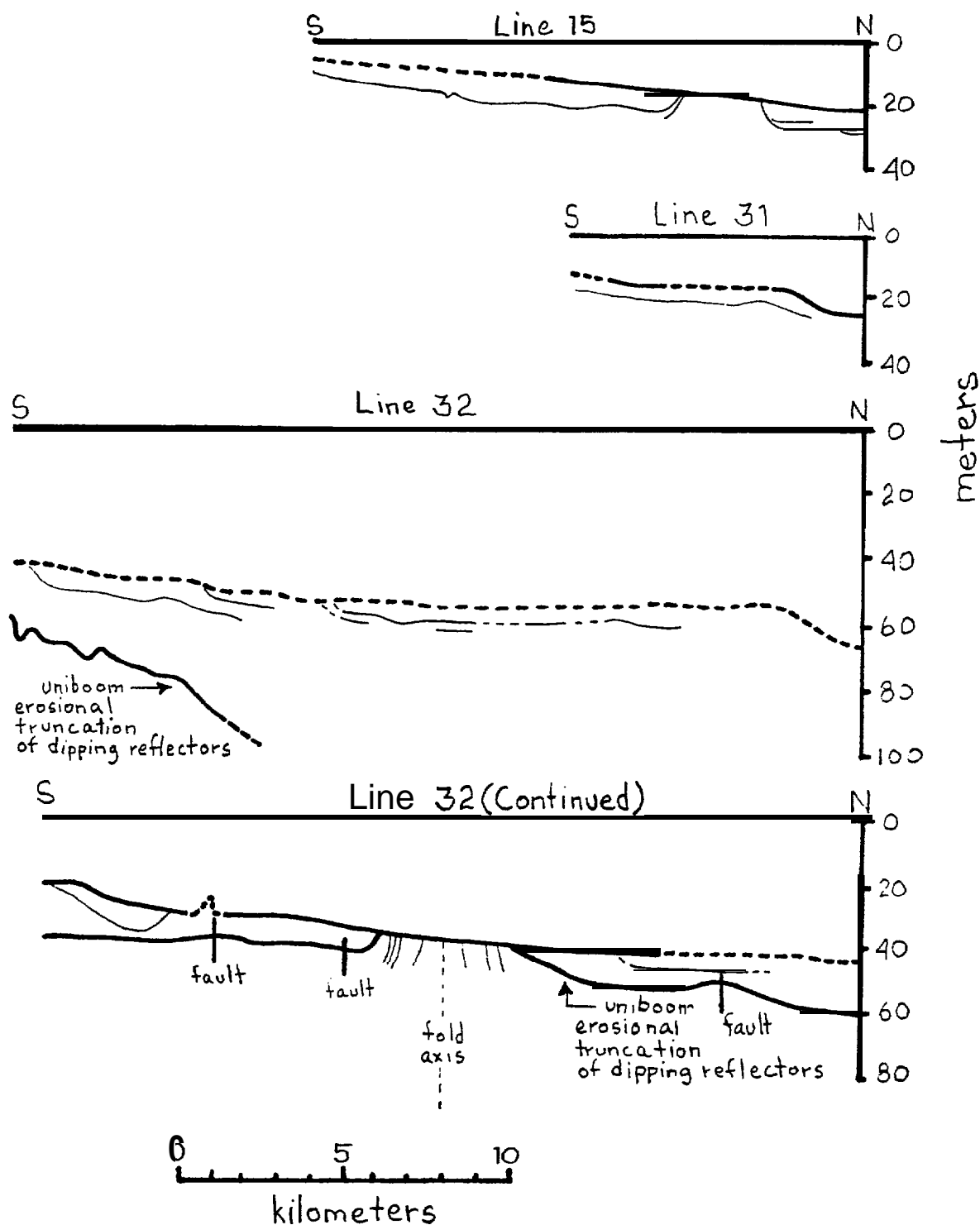


Figure 20. - Line drawings of 7kHz and **Uniboom** sub-bottom reflectors from **tracklines** between Camden Bay and Barter Island. Surface sediment textures in Figures 20 thru 23 are interpreted from sonargraphs and fathograms.

Surface Sediments : Cohesive ——— Non-cohesive - - - - -

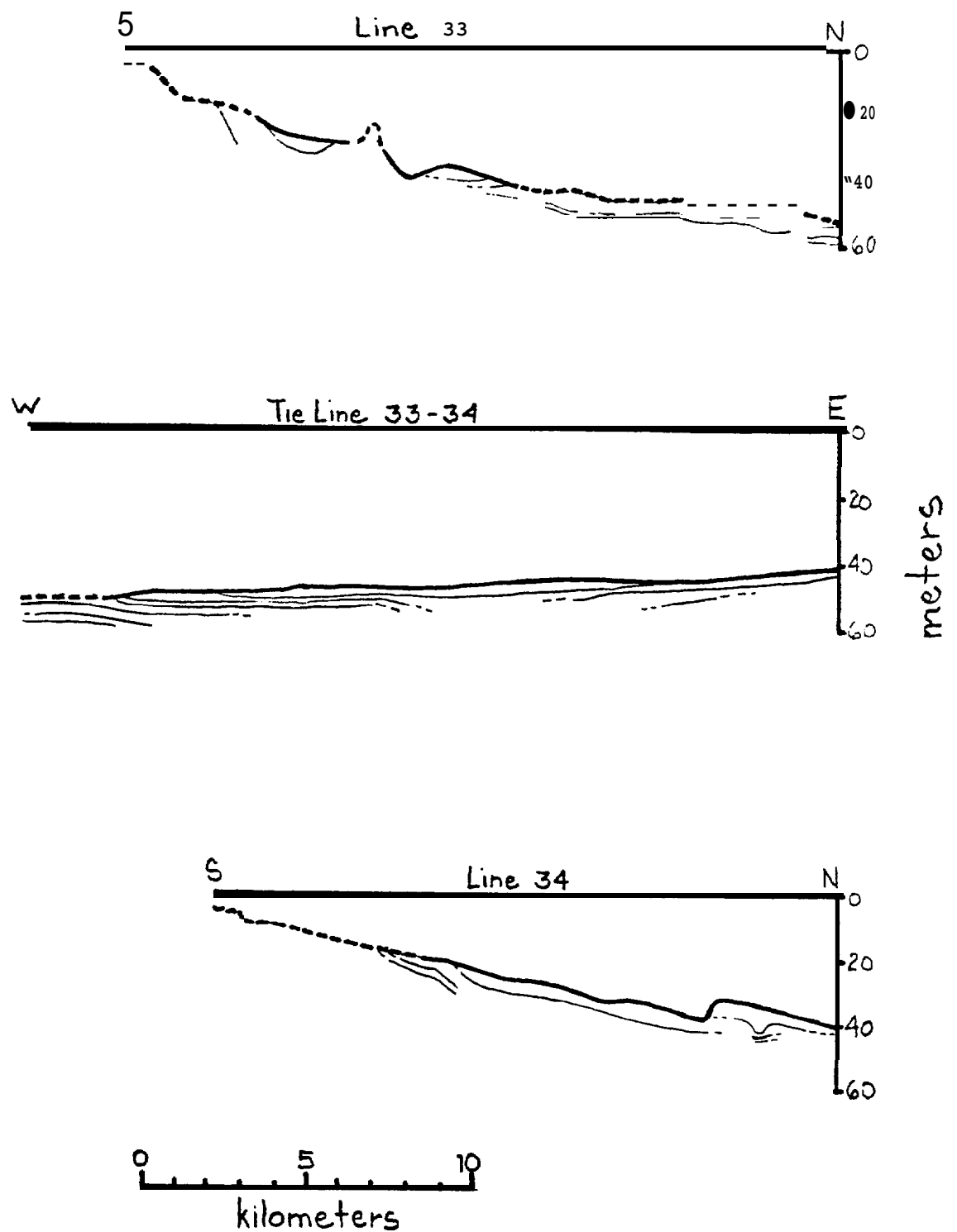


Figure 21.- Line drawings of 7kHz sub-bottom reflectors from **tracklines** between Barter Island and the Jago River.

Surface Sediments: Cohesive ——— Non-cohesive - - - - -

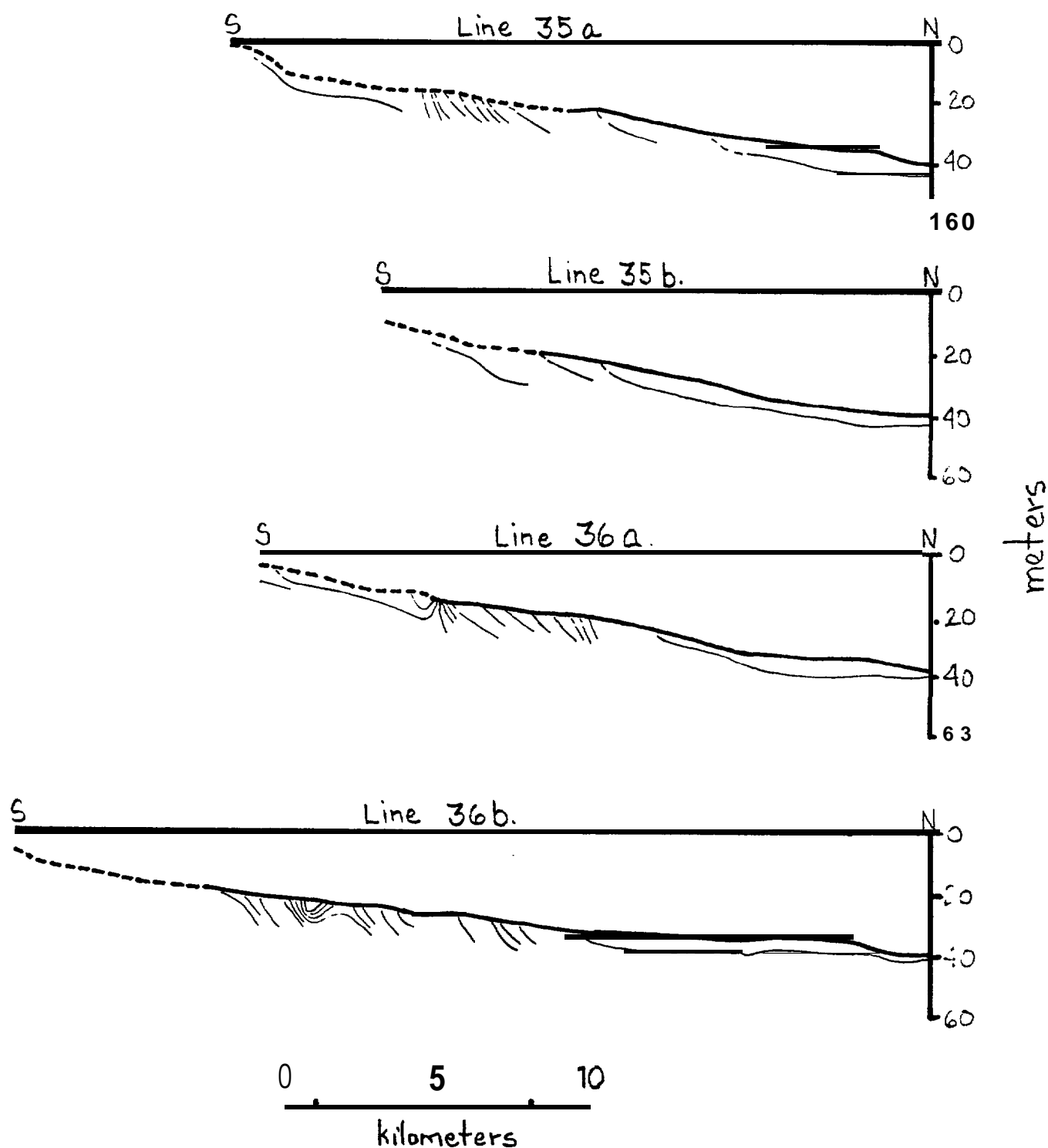


Figure 22. - Line drawings of 7kHz sub-bottom reflectors from track lines between the Jago River and Beaufort Lagoon.

Surface Sediments : Cohesive — Non-cohesive - - - - -

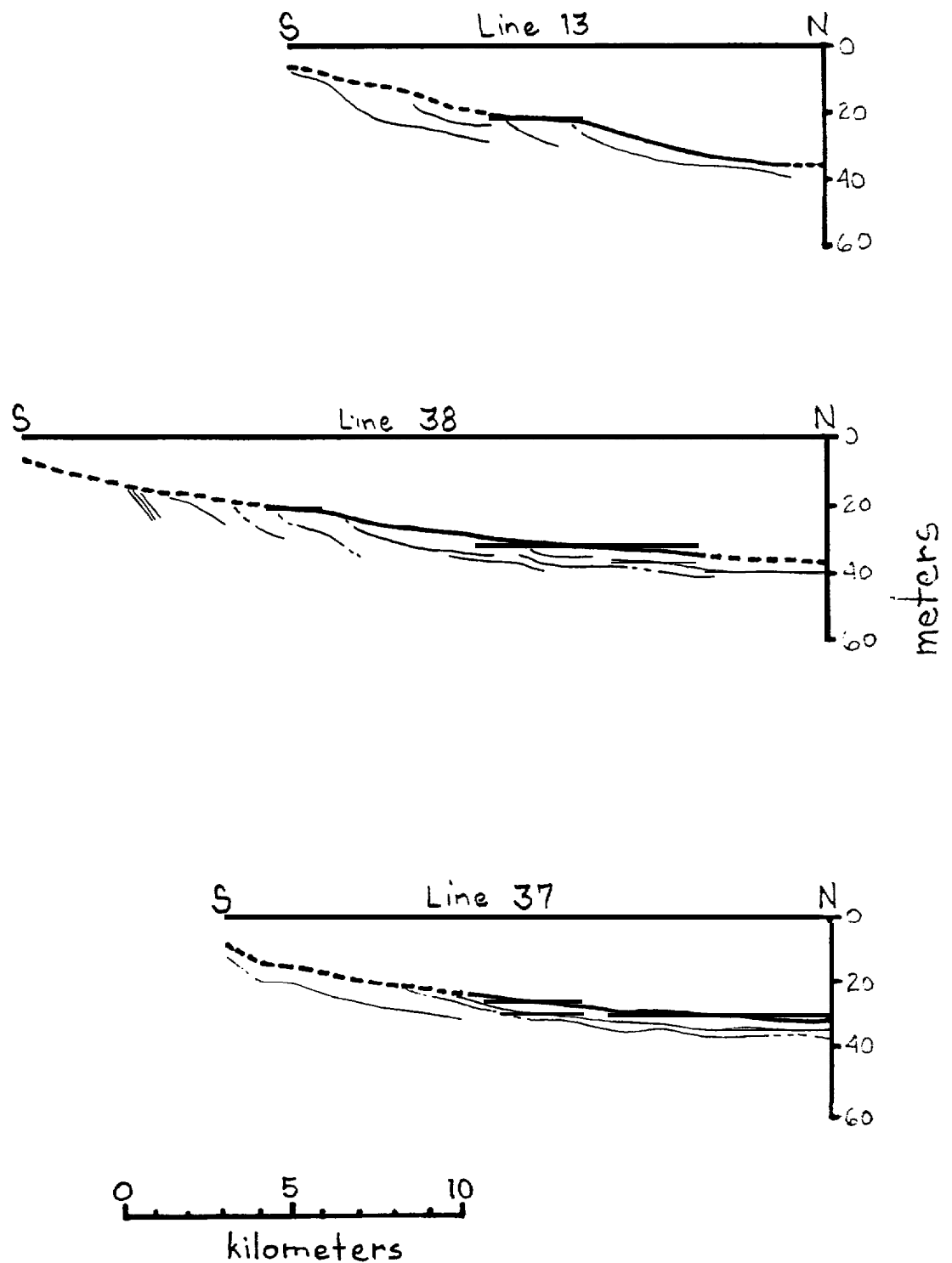


Figure 23. - Line drawings of 7kHz sub-bottom reflectors from **tracklines** between Beaufort Lagoon and Canadian border.

The seafloor trace also distinguishes between surface material types, as interpreted in the next section.

None of the sections traced in figures 20 through 23 contain reflector patterns revealing sediment accretion. On the contrary, most areas show shallow **subbottom** reflectors at varying angles to the sea floor, and cropping out somewhere along the traverse. We can detect no thickening of surface units towards rivers and coastal bluffs, the modern sediment sources. The tracings also do not reveal a thickening of units towards the shelf edge. Much more **work** will be necessary to gain an understanding of the stratigraphic complexities below the shelf surface. We prepared a **scattergram** with water depth plotted against sediment thickness above the first reflector (Fig. 24) and found that in the areas covered by our **tracklines**, the first reflector thickness is nowhere greater than 10 m and in most cases is less than 6 m. **Surficial Sediments** - In our appraisal of surface sediment textures for the region from the Canning River to the Canadian border we used the surface sediment samples collected in 1981, the classification of geophysical records into cohesive and non-cohesive sediment types in 1-km-track segments, and sediment **analyses** of samples reported by P. W. Barnes (1974).

The 1981 shipboard sample descriptions are condensed in table 2. Dots mark the sampling sites in figure 25 (station numbers are shown in Fig. 2). The comparison of the texture of surface sediment samples with the appearance of ice gouge relief on fathograms and monographs showed good correlation. Our interpretation of the geophysical records and the classification of relief forms into "rough" and "subdued," and classification of surface sediment textures into "cohesive" and "non-cohesive," is, of course, strongly influenced by detailed diving and sampling investigations made west of the Canning River. Figure 19 is a sample of **fathograms** recorded in areas of cohesive, muddy surface sediments (A) and non-cohesive, coarse, granular sediments (B). In the latter case the materials piled up in flanking ridges during the ice-gouging process move **downslope** to assume the angle of repose as the ice passes. Subsequently the aging process, aided by current effects on non-cohesive materials results in broadly rounded ice-gouge forms. The **fine-grained** surface sediments, on the other hand, assume relatively steep slopes, **sometimes** blocky shapes, during disruption by ice and remain in this position even through periods of current activity. The monographs shown in figure 26 represent samples of these **two** distinct bottom types. In figure 26A the gouges are cut into cohesive materials, most apparent in the ridge details. These are piles of jagged materials alined along gouges and lack the continuous smooth ridges seen **in** figure 26B. The smooth ridges of figure 26B were recorded **at** the same time and place **as** Fig. 19B. Figure 26A was chosen as an example because a first-year pressure ridge that produced the rake marks on the seafloor is firmly grounded at the end of the gouges.

The **two** bottom types interpreted from the geophysical records were plotted and the results are shown on the map in Figure 25. Coarse, granular materials blanket a strip from the coast to about 15-m water depth. Seaward of the 15-m water depth lies a zone of fine, cohesive surface sediments, which grade seaward into coarse granular materials. Coarse-grained materials can be traced uniformly for many kilometers on line 32, the long track extending northwestward from Barter Island to the shelf break. **At** 53-m depth we

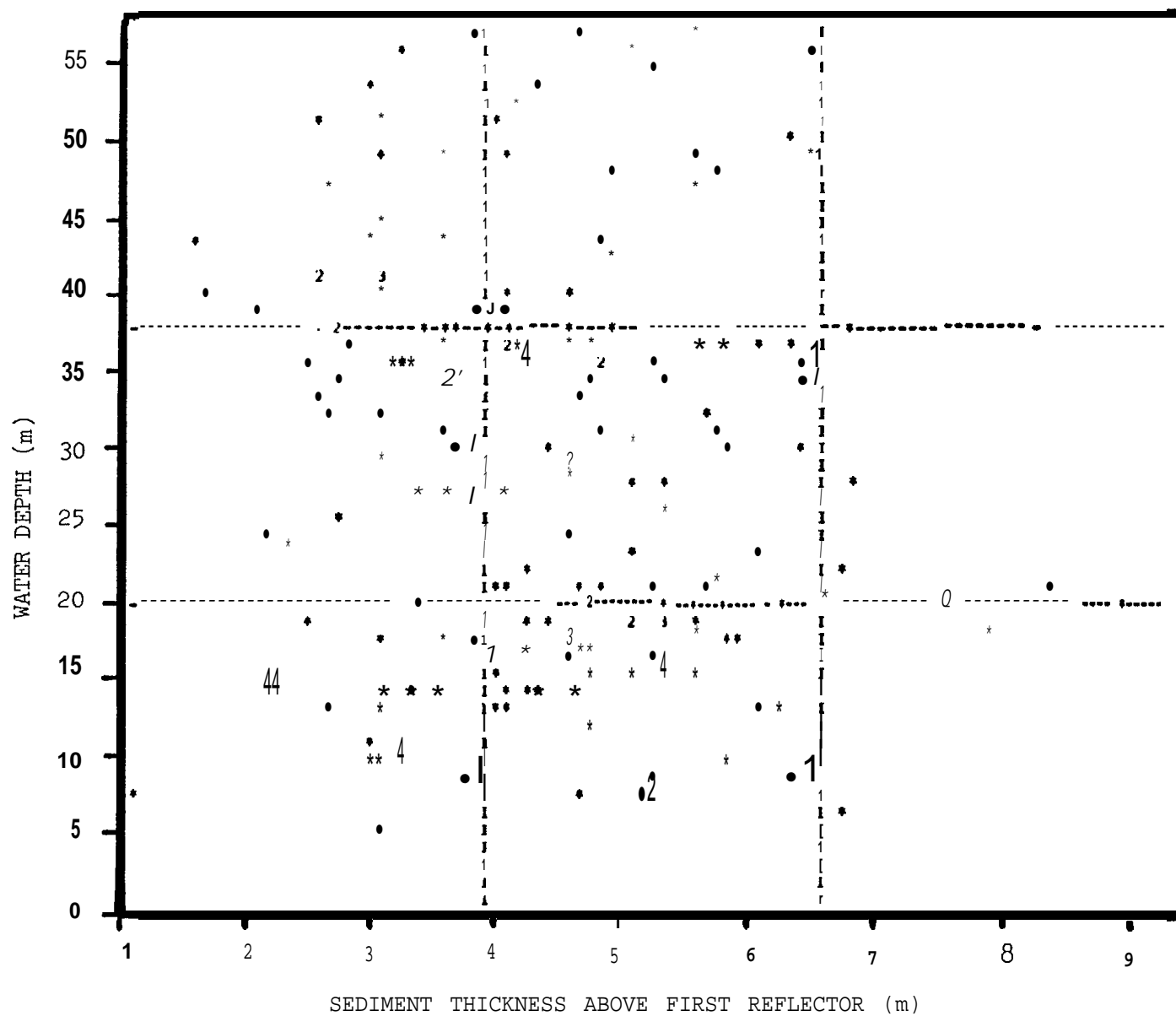


Figure 24. - Scattergram of sediment thickness above the first reflector versus water depth.

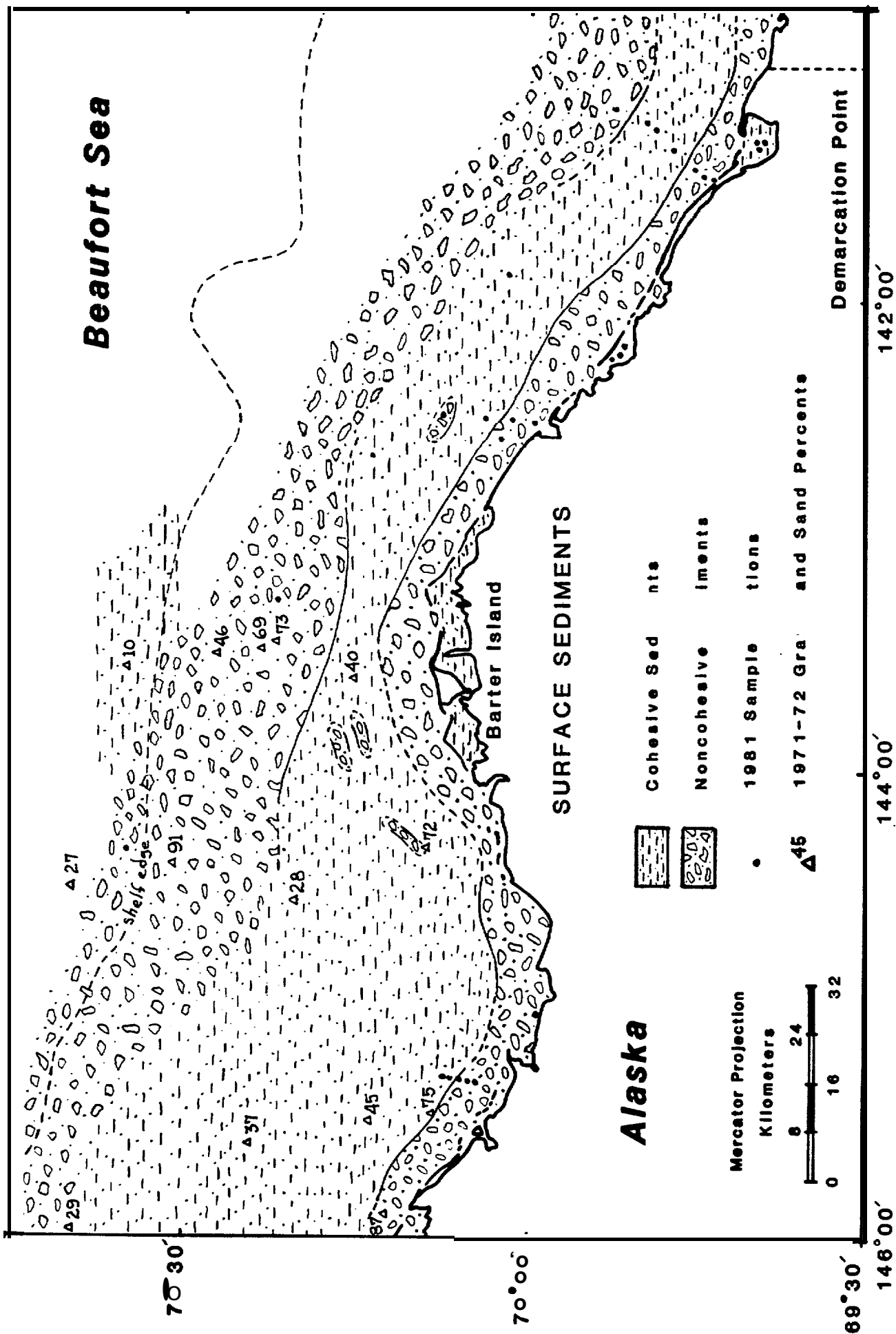


Figure 25.- Map of surface sediment textures, as interpreted from geophysical records and sediment samples. Percentages of combined sand and gravel are next to the old sample stations of Barnes (1974).

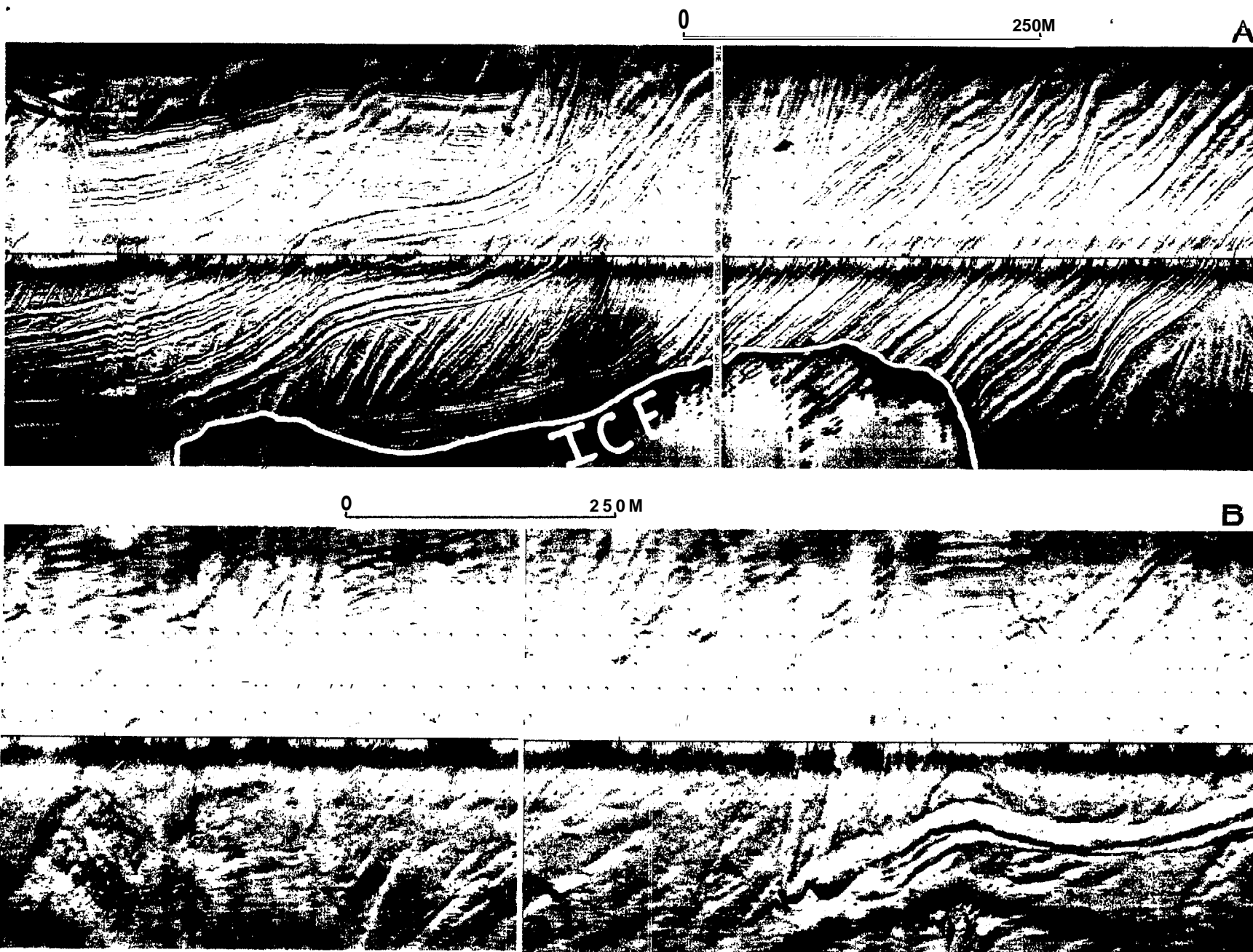


Figure 26. - Sonographs of rough (A) and smooth (B) gouge relief, a difference we interpret as due to the presence of cohesive (A) and non-cohesive (B) sediments. A large piece of ice is grounded in the lower center part of sonograph A; at gouges under the ice are hidden. Sonograph B corresponds to fathogram B in Fig. 19. The 3 sharp gouges on the fathogram are clearly visible on the sonograph. Even though there is a lot of relief in the area, the seafloor texture is smooth because of the non-cohesive nature of the sediments.

interrupted the line to collect a sample for verification and retrieved essentially clean gravel with attached organisms. The shoals within the strip of cohesive materials on the central shelf appear to be generally sand and gravel. The numbers shown on the shelf west of Barter Island in figure 25 represent percentages of sand plus gravel taken from surface sediments analyzed by Barnes (1974). These values substantiate that much of the shelf surface, and especially the outer half, is covered with coarse granular materials.

Shoals of the stamukhi zone

The relationship of coastal promontories and shoals acting as strong points in the control of ice dynamics and zonation has been of considerable interest to our studies (Rearic and Barnes, 1980; Reimnitz et al. , 1978). The published charts for the study area do not show a pattern of shoals down-drift of the Barter Island promontory, similar to the pattern developed west of the Cross Island promontory. However our reconnaissance survey lines provide single crossings of a number of shoals. One long linear shoal off the Canning River was crudely defined by a number of crossings. A number of samples collected around that shoal show it to be composed of sand and gravel, similar to the shoals west of the Canning River which have been thoroughly studied. Most of the other shoals as well are composed of coarse granular materials as interpreted from the geophysical records. A sample crossing is shown in figure 27. The sonograph shows an intensely gouged sea floor on both sides of the shoal. Here the gouge flanks have the rough appearance typical of flanks associated with fine-grained cohesive materials. The shoal itself, is composed of coarse granular material with a smoothed, rounded surface and a trace of current ripples on the crest. Ice hangups are most common on such shoals and the sonograph shows such a stamukhi along the crest.

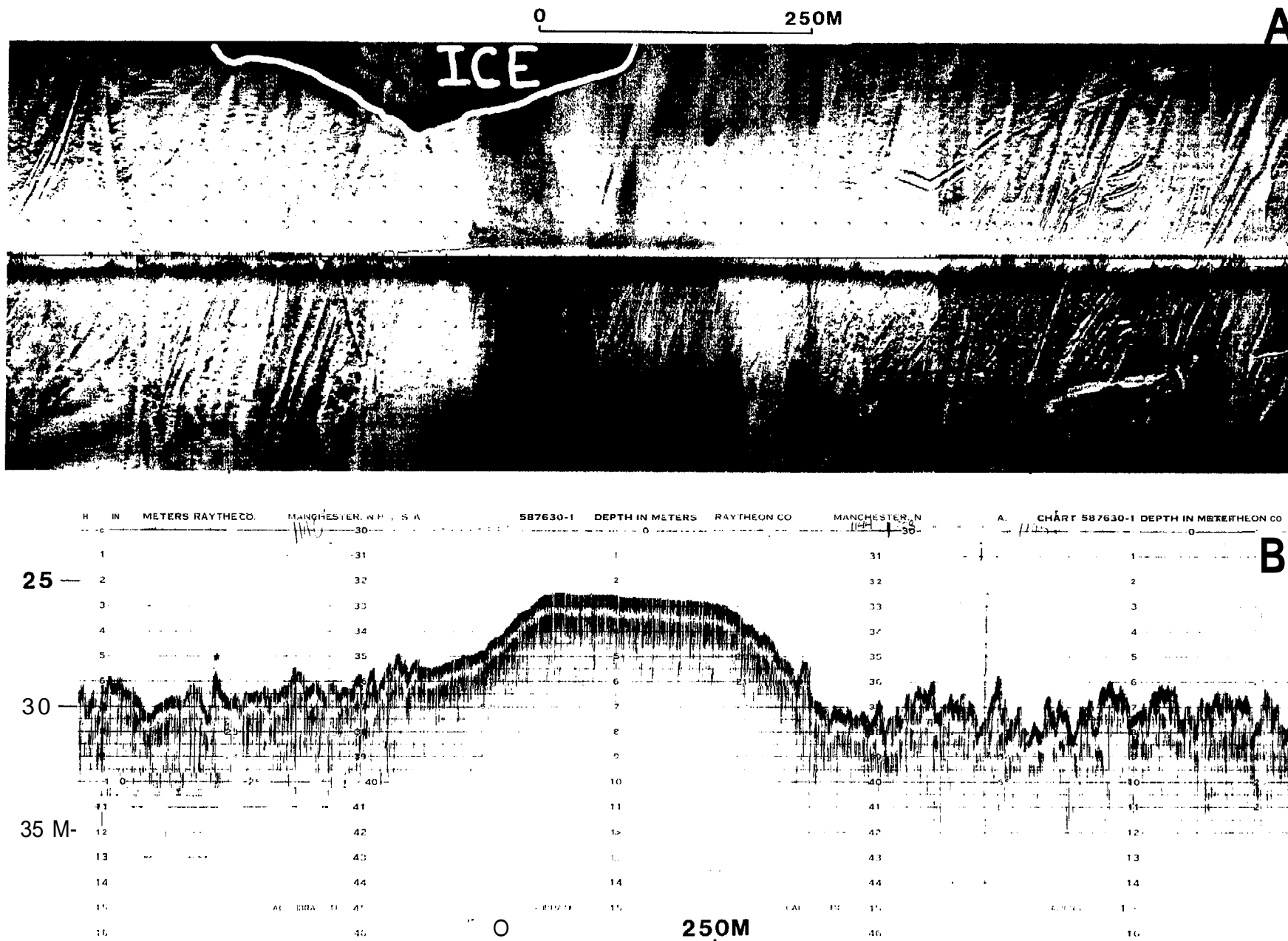


Figure 27. - Sonograph (A) and fathogram (B) of a shoal crossing in the stamukhi zone. The shoal crest is characterized by a smooth, noncohesive texture, while the surrounding bottom exemplifies the jagged, cohesive sediment texture. The smoothness of the shoal crest, which is gouged by ice more frequently than the surrounding bottom, indicates that hydraulic processes rapidly rework the sediments on the shoal crest.

DISCUSSION AND CONCLUSIONS

Sedimentation - From combined coring, diving, and seismic profiling studies in numerous different marine geological environments and settings west of the Canning River, we are convinced that where repeated ice plowing occurs with slow sediment accretion, no continuous sedimentary units develop. Sediments come to rest mainly in troughs of gouges, and the shape and extent of a trough define the limits of sedimentary units. Assuming that the **depositional environment** on the Beaufort Sea shelf has remained constant for the last 10,000 years or longer, with slow accretion and rapid **re-gouging**, we see no possibility for the blanket of Holocene sediment to contain continuous internal reflectors.

All indications are that modern sediment accumulations, possibly present in lagoons and bays, are essentially lacking on the open shelf. The fine grained, cohesive sediment mapped in a band on the central shelf, may be modern deposits of several meters thickness, and most likely the shoals of the stamukhi zone are constructional features post-dating the last transgression. The coarse granular materials on the inner shelf and on the outer shelf seem to be relict deposits. The relict nature of the shelf edge gravels has been discussed by Barnes and Reimnitz (1974), **Naidu** and Mowatt (1974), and Rodeick (1975). Their interpretations are based on a) low rates of modern ice rafting of coarse **clasts** compared to overall sediment accretion rate, b) observed ferromanganese coatings on cobbles, c) about 15,000 year old **C¹⁴** ages for near-surface shelf edge and upper slope sediments, d) source rock considerations, and e) lack of seaward decrease in sediment grain size from coarse grained near the sediment source to fine grained near the outer edge of the shelf.

Grantz and Dinter (1980) mapped a seaward thickening wedge of Holocene silt and clay on the Beaufort Sea shelf, using high resolution seismic reflection records. In the Barter Island area in particular, they show a large area of structurally formed and truncated stratified deposits lacking any Holocene marine sediments, and flanked on the northeast- and northwest side by Holocene marine sediments thickening to 30 and 40 m at the shelf edge. Line #32 of the present study was aimed at reaching the shelf edge where modern marine sediments are 40 m thick, where sedimentation rates presumably are high, and where the greatest water depth at which ice gouges exist would correspond with the present maximum ice keel depth to be encountered within the Beaufort Gyre. We reasoned that rapid sediment accretion suggested by 40 m of modern sediment **would** eliminate gouges within a period of several hundred years. Line #32 (for cross section see figure 20) does indeed cross the erosional region on the mid shelf, **where** older sediments are truncated by the seafloor, but it does not show a thick homogeneous wedge of Holocene sediments to seaward. The character of the gouges recorded, in fact, made us suspect gravelly surface sediments and we interrupted the line to collect a sample. The gravel retrieved at 52 m water depth, along with the homogeneous appearance of the records for tens of kilometers, supports previous **sedimentological** interpretations that much of the outer shelf in the eastern Alaskan Beaufort Sea is blanketed by relict gravels, and not by Holocene marine sediments.

One of the major potential modern sediment sources for the eastern Alaska

Beaufort Sea shelf is the Mackenzie River. Therefore, a comparison with the sediment distribution on the shelf between our study area and the Mackenzie Delta will shed additional light on our belief that the outer shelf off northern Alaska is a surface **of** non-deposition. Figure 28 is a compilation of our sediment texture map extending to the Canadian border, and a map of **sand-plus-gravel** percentages for the region east of the border by Vilks et al. (1979). The Canadian shelf surface is covered by sand and gravel. Yorath et al. (1970) interpreted the sandy gravels, sands, and hard pebbly **lutites** as "relict glacial deposits and ice-pressed tills." Thus, these combined interpretations of shelf surface sediments, while not matching across the border in detail, leave no room for a thick wedge of Holocene silt and clay on the outer shelf. Our interpretation of the 1981 seismic data also rules out the possibility of thick Holocene sediments in our study area. A thorough study of this problem is urgent because the interpretation that slumping, sliding, and faulting are active geohazards in this area (Grantz and Dinter, 1980) is strongly dependent on whether the shelf edge sediments are old or recent.

Ice Gouging - The statistical mean values calculated for various ice gouge parameters in the present study area are greater than those of the area west of the Canning River (Barnes et al., 1982). This can be explained by the exclusion of surveys in lagoons and bays from our present data analysis. Aside of this difference, the overall patterns are found extending all the way to the Canadian border and probably beyond. Along the entire Alaskan shelf, the 18 m isobath separates inshore low density and size values from offshore high density and size values. The **stamukhi** zone, lying between 18 and 36 meters of water depth, in all areas stands out by having the highest values on most parameters measured, but east of the Canning River the values do not decrease offshore with the same consistency as to the west of the Canning River. Gouge densities follow the most consistent pattern **along** the entire shelf. In the present study area, the pattern of highest gouge densities corresponds rather well with a 5-year composite of ice-ridges prepared by Stringer (1978) and shown in figure 29. The significance of the 13 m **isobath** as a boundary between areas of mild and severe ice hazards (Kovacs, 1980) has not shown up in our data analysis for the length of the shelf.

The trends of water depth contours in the present study area are more northerly on the average than those west of the Canning River, and a comparison of ice gouge trends in the two regions supports previous conclusions that the plowing action aligns with the **isobaths**. In this study we were again able to demonstrate the tendency for ice gouges to align more consistently **isobath-parallel** on the up-drift (eastern) side of major promontories, and more variably on the down-drift side (Barnes, et. al. , 1982).

The lack of gouges on the crests of shoals in the **stamukhi** zone, and the presence of hydraulic **bedforms** in coarse granular materials, again supports our contention that active hydraulic processes reshape, and perhaps help to rebuild, features that should soon be eliminated by ice scouring. Even in the consistent presence of **stamukhi** (grounded floes) on the shoals during surveys (figure 27) we rarely detect gouges, while the surrounding low and more protected terrain with cohesive surface sediments is highly gouged.

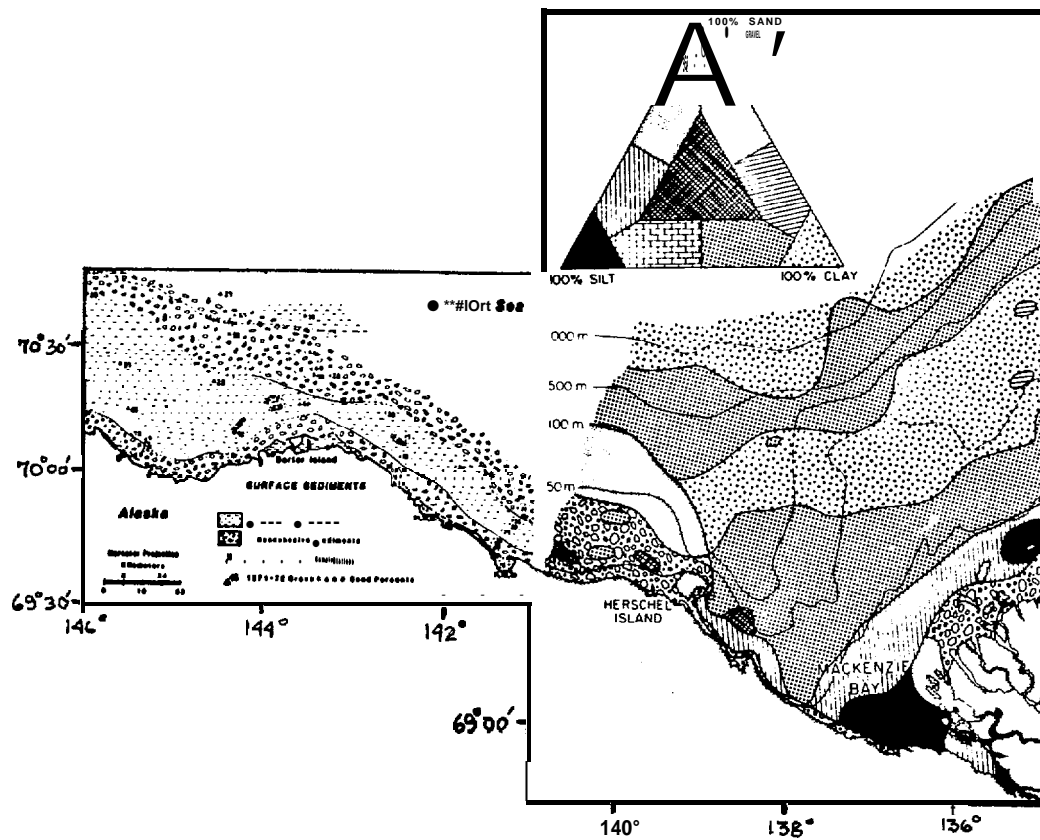


Figure 28. - Composite of surface sediment textures from the present study, and east the Canadian border (Vilks et al. , 1979).

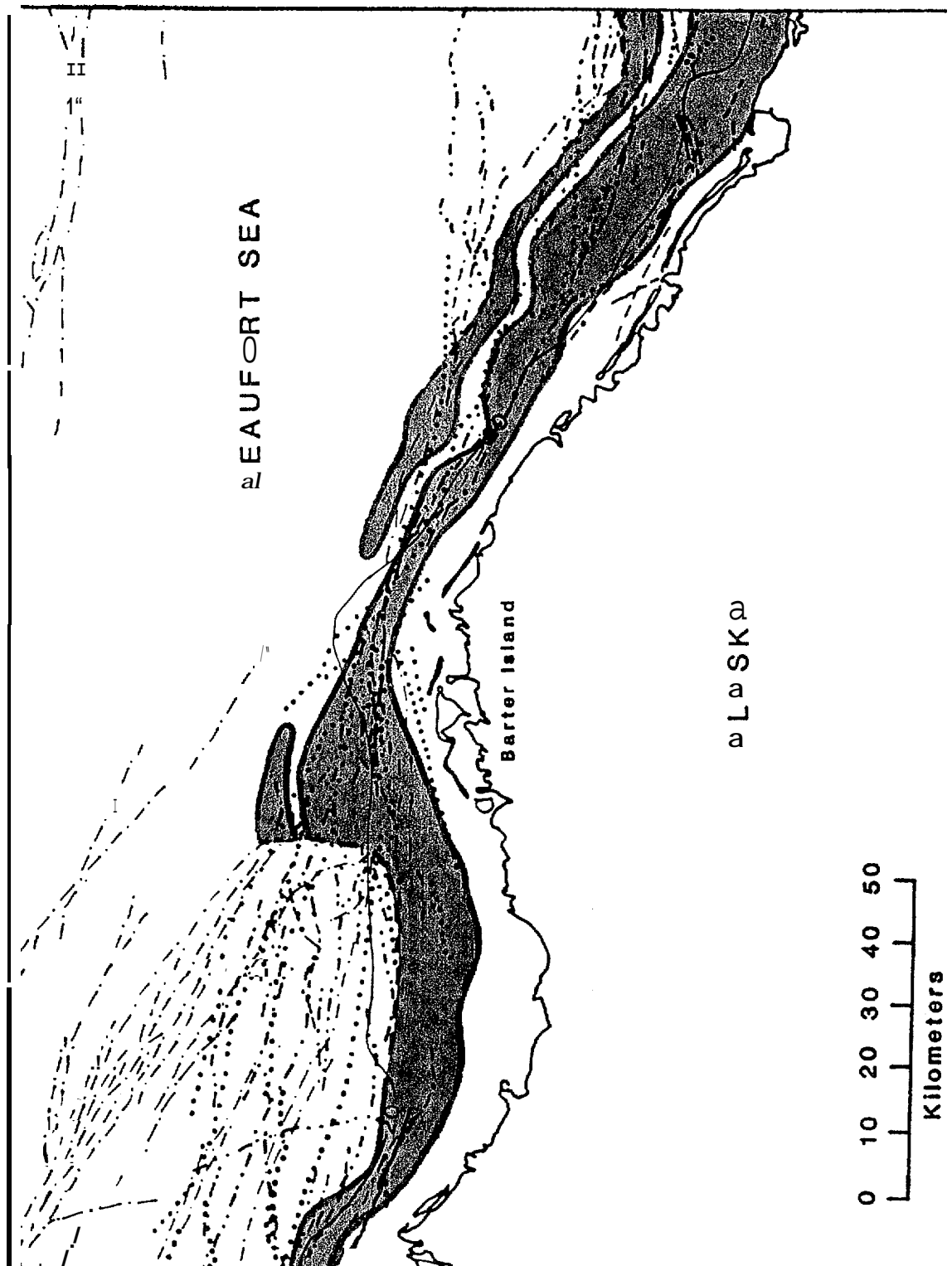


Figure 29.- A composite of ice ridges in the Barter Island area from 1973 to 1977, from Stringer et al., 1978. The shaded area indicates zones of high density gouging (> 100 gouges per km) that were contoured from data of the present study (see figure 6).

The total vertical relief possible for a single gouge was previously estimated (Barnes, et al., 1982) by adding the highest ridge from one gouge to the deepest trough of another. In the recent surveys we found 8 meters **total** relief across a single gouge, leading us to believe that accurate estimates of ice gouge extremes can now be made from our large volume of data.

Drifting ice scraping the seafloor appears to be an efficient **planation** agent, producing erosional unconformities and truncating thick sets of dipping strata. We feel that hydraulic processes alone acting on that same surface would have sculptured it in accordance with the resistance to erosion offered by the different geologic units. Relatively well indurated beds **would** form scarps. The ice pack acting on an extensive, non-homogenous surface, however, seems to take the different **lithologic** units down to the same level by focusing mainly on the high points. Viewed in this light, the existence of major, well defined shoals, is more perplexing.

So far we have been unable to relate the intensity of ice gouging to the underlying geology. Thus, one could also argue that all geologic strata exposed to the action of ice in the study area are weak compared to the forces of the **moving** ice keels.

New Evidence for greater than expected ice depth - Favorable ice conditions in 1981, and a relatively narrow shelf east of the Canning River enabled the R/V **KARLUK** to survey ice gouges in generally greater water depth than has been possible in the western sector. One particular line was extended to the very shelf edge. In general, the relationship between ice gouging and water depth in the study area is similar to that determined for areas west of the Canning River, with lowest values for certain gouge parameters inshore and offshore of the stamukhi zone. In the present study, ice gouges were traced to maximum water depth of 58 meters. Beyond that we saw only very broad, subdued relief features unrelated to ice keel interaction. Among the **bedforms** beyond the deepwater gouge limits we found slope-parallel, rhythmic lineations of 3 m wave length but less than 20 cm of relief, which we interpret as probable hydraulic **bedforms**. These indicators, along with the presence of surface gravels rather than fine materials, the sub-clued nature of gouge relief forms, the seaward decrease in ridge height relative to trough depth and width, and especially the recorded current pulses of up to 50 cm per second along the shelf edge (**Aagard**, 1977) all suggest that active currents rework the deep water gouges. Based on these considerations, the gouges found at 58 m water depth are modern rather than relict (produced during lower stands of sea level). Surficial hyperbolic reflections on **Uniboom** crossings of the shelf edge between Barrow and the Canadian border, and the accompanying surface roughness, are fairly certain indicators for the presence of ice gouges. These indicators can be traced in 28 representative traverses to maximum water depths of between 60 and 64 meters (Dave **Dinter**, U.S. Geological Survey, oral communications, 1982).

Our previous contention that ice gouges seen on the Beaufort Sea shelf at depths greater than 47 m (the deepest keel actually observed) are modern has recently found additional support. Marine geologic studies by Canadian **workers** in the Southern Beaufort Sea no longer call for lower sea levels to account for the deepest gouges observed. **Also**, statistical treatment of ice keel distributions in Arctic deep water, allow for 60 m deep keels to occur at

arate of one every few hundred years (Peter **Wadhams**, oral communication, 1982). These findings are of little consequence at the present stage of petroleum development in the Alaskan Beaufort Sea, but may in the future assume considerable importance.

Shallow seismic stratigraphy - Our analysis of seismic records has not progressed to the stage where correlating individual units from line to line, and their interpretation, can be attempted. However, we can put some limits on the surface units - the Holocene marine sediments. Our reasoning leading to the conclusion that Holocene marine sediments cannot contain continuous seismic reflectors has been presented above. This is not only a theory, it has been proven true in numerous site specific studies in the west. Based on this fact, the sediment thicknesses above the first sub-bottom reflector are the upper limit for the thickness of Holocene marine sediments. A plot of these values (Fig. 24) against water depth shows no trend. The mean depth below the sea floor is **nearly** 7 meters. But as discussed before, the geometry of units defined by the shallow reflectors, excludes them in most regions from being Holocene marine sediments. They are in fact older units.

Thick sections of stratified, tectonically deformed, probably Pleistocene strata dipping at various angles, are truncated by the seafloor over extensive regions in the Barter Island area. We have not been able to trace any portions of the section to Barter Island from the **Flaxman** Island area, where well known stratigraphy exists from **boreholes**. Some faults extend to near the sea floor, but we are unable to detect surface scarps or other signs of recent fault displacements. However, the smooth truncation surface, extending for many kilometers cutting across numerous **strata** of presumably different **erodability**, suggests that ice scouring is an efficient **planation** agent that treats all **materials** available uniformly. **Thus**, the lack of modern fault scarps in our data is not necessarily evidence against recent movement postulated by Grantz and **Dinter** (1980).

Sand and Gravel Resources - Triggered in part by the high demand for sand and gravel as construction material for offshore petroleum development, the Federal Government is making preparations for **managing** these resources on the Arctic shelf through a leasing program. In the present study area all indications are that gravel is plentiful, even in deep water, and need not be hauled great distances. In areas where active gouging creates up to 8 m of vertical relief, the seafloor reflectivity and overall appearance is homogeneous for many kilometers. If such areas on the outer shelf were underlain by interbedded mud, sand and gravel, the plowed ridges **would** reveal such inhomogeneities. The sea floor **would** be littered with slabs of stiff silty clay. The appearance of the geophysical records suggests to us that on the outer shelf fairly clean, **coarse** granular materials have a thickness of at least several meters. However, several box cores from the outer shelf contain firm mud units (Barnes and **Reimnitz**, 1974), raising questions that need answers.

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APPENDIX

to

ATTACHMENT "I"

T = transitional S = smooth R = Rough

C = cohesive N = non-cohesive D.M.P.

Line Number : 08 Year : 1981									
Fathogram Measurements					Sonograph Measurements				REMARKS
SEGMENT	Water Depth (meters)	Reflector "A"	Gauge Depth	Ridge Height	Gauge Width	Density C _D	Orientation (°T)	Sediment Cohesion	
000	24.3	—	.4	.4	17	168/161	105 / 285	R/C	4 Km of sonar missing. Count adjusted. ^{SSS} Range 150 m
001	22.8	—	.8	.9	—	—	—	R/C	no sonar for this segment.
002	22.1	27.1	.8	.6	—	—	—	R/C	no sonar for this segment.
003	21.3	25.5	.4	.5	—	—	—	R/C	no Sonar for this segment
004	17.4	22.7	<.2	<.2	—	—	—	S/N	no sonar for this segment. 4 meter bench at start of segment (includes 1 meter shoal on crest).
005	17.0	22.0	.2	.3	3	97/93	080 / 260	S/N	4 Km of sonar missing. Count adjusted.
006	16.1	20.6	<.2	<.2	11	88/127	156 / 246	S/N	1 multi-gauge - 12 inc., 98 m, 144°T
007	17.3	22.3	<.2	<.2	5	70/112	20 / 110	S/N	gauge orientations are highly variable.
008	18.0	23.5	<.2	.2	5	87/106	36 / 126	S/N	" " " " "
009	18.8	24.3	.2	.2	4	142/227	158 / 248	R/C	Orientation is highly variable. 3 Multi-Gauges - 24 incs. in largest - 125 m wide - 240°T
010	19.1	24.4	.4	.3	8	152/219	24 / 114	R/C	
011	19.6	25.3	.4	.5	7	175/231	151 / 241	R/C	orientation is highly variable.
012	20.1	25.3	.5	.7	5	169/243	26 / 116	R/C	orientation is highly variable. large multi gauge with 13 incisions and oriented 90°T. width = 65 m.
013	20.2	25.8	.4	.5	5	124/119	70 / 250	R/C	sonar quality poor - many range changes.
014	18.1	24.5	.4	.5	16	156/190	143 / 323	R/C	orientation is highly variable.
015	16.9	22.8	.3	.5	5	121/174	26 / 206	R/C	600 m of sonar missing. Count adjusted.
016	15.9	21.4	<.2	<.2	5	39/48	126 / 306	S/N	4 Meter shoal in middle of segment. most gauging ends on offshore slope of the shoal.
017	12.7	18.2	<.2	<.2	4	25/46	166 / 256	S/N	
018	13.0	—	.2	.2	10	42/51	35 / 125	S/N	
019	15.8	—	.2	.3	7	137/167	145 / 235	T	orientation is highly variable.
020	16.4	—	.2	.3	5	149/197	150 / 240	R/C	orientation is highly variable.
021	17.0	—	.4	.9	—	—	—	R/C	no sonar this segment
022	16.9	—	.3	.6	6	167/240	23 / 113	R/C	orientation is highly variable.
023	17.8	—	.5	.5	5	148/213	23 / 113	R/C	orientation is highly variable.
024	17.5	22.8	.9	.9	18	172/187	136 / 226	R/C	orientation is highly variable.
025	18.2	24.9	.6	.6	7	136/148	45 / 225	R/C	orientation is variable.
026	17.2	22.5	.7	.5	—	—	—	R/C	no sonar this segment.
027	16.4	20.9	.4	.4	4	85/83	120 / 300	T	400 m. of sonar missing. Count adjusted.
028	13.7	19.0	.2	<.2	3	16/17	128 / 308	S/N	3.5 m. shoal in middle of segment

1992

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gauge measurements in

Line Number : 09

Year : 1981

gouge measurements in meters

SEGMENT	Fathogram Measurements				Sonograph Measurements				REMARKS
	Water Depth (meters)	Reflector "A"	Gouge Depth	Ridge Height	Gouge Width	Density	Orientation (° T)	Sediment Cohesion	
029	12.1	16.5	.3	.3	10	115/131	15/90	R/C	
030	11.9	16.5	.5	1.0	17	69/144	8/83	R/C	
031	12.2	16.2	.5	.9	3	108/104	113/113	R/C	
032	13.3	16.6	.2	.8	7	105/103	60/80	R/C	
033	14.0	17.9	.5	.7	8	171/168	61/81	R/C	Very high gouge density but gouges are small and shallow.
034	14.8	20.0	.5	.6	5	217/213	58/78	R/C	gouges are narrow (~1-2m wide)
035	15.4	19.9	.3	.7	8	263/214	52/72	R/C	400m of sonar missing. Count adjusted
036	15.8	20.4	.3	.6	3	165/172	52/72	R/C	high density gouging stops at middle of segment for no apparent reason (no shoals, etc.)
037	16.1	19.9	<.2	.3	7	58/59	57/77	R/C?	
038	16.5	22.3	<.2	.2	3	50/61	153/173	R/C?	(Because they are shallow they may be covered by a transitory layer ??)
039	17.2	—	<.2	.2	7	56/58	51/71	R/C?	
040	17.9	22.1	.3	.5	—	—	—	—	NO SONAR THIS SEGMENT. HIGH DENSITY GOUGING STARTS AGAIN AT SEG. START (NO APPARENT REASON).
041	19.0	23.5	.3	.7	8	226/362	18/38	R/C	2 SUB BOTTOM REFLECTORS - TOP ONE APPEARS TO SURFACE AT ~40 Km point.
042	19.7	24.3	.3	.3	5	140/141	53/73	S/N?	long low shoal with little gouging on it. Probably sand and will not hold gouges. 500m Sonar missing-adj.
043	19.0	—	.4	.6	7	214/216	123/238	R/C	high density gouging starts again after shoal crossing.
044	20.0	24.0	.3	.5	10	198/200	126/241	R/C	END OF SONAR
045	18.5	23.0	.3	.4	—	—	—	—	
046	18.2	22.3	—	—	—	—	—	—	2 meter shoal at beginning of segment
047	18.6	22.1	—	—	—	—	—	—	
048	17.2	—	—	—	—	—	—	—	
049	15.3	—	—	—	—	—	—	—	
050	14.5	—	—	—	—	—	—	—	2 meter shoal at mid segment.
051	13.9	—	—	—	—	—	—	—	
052	11.6	—	—	—	—	—	—	—	
053	12.1	—	—	—	—	—	—	—	
054	11.2	—	—	—	—	—	—	—	
055	9.5	—	—	—	—	—	—	—	

[illegible]

Line number : 15

Year : 1981

gouge measurements in meters

[illegible]

[illegible]

Line Number : 32									
Year : 1981									
gauge measurements in meters									
Fathogram Measurements					Sonograph Measurements				REMARKS
SEGMENT	Water Depth (meters)	Reflector "A"	Gauge Depth	Ridge Height	Gauge Width	Density	Orientation (°T)	Sediment Cohesion	
000	65.6	—	—	—	—	0/0	—	N	Possible sand waves. Very subdued. $\theta = 116^\circ T$ Range 150m. $h = < 20 \text{ cm}$ $\lambda = 3 \text{ m}$
001	64.2	—	—	—	—	0/0	—	N	Possible sand waves and ripples. $\theta = 118^\circ T$ $h = < 20 \text{ cm}$ $\lambda = 3 \text{ m}$
002	57.3	—	1.5	.7	11	12/13	137/102	S/N	Range change to 200m. Possible shelf edge at midsegment. Probably granular sediments.
003	53.6	—	1.1	.8	13	42/44	131/96	S/N	Gauges are large but subdued on sonar (sand & gravel). Most gauges are wide ($> 5 \text{ m}$).
004	54.3	—	.8	.5	8	41/41	125/90	S/N	
005	54.2	—	1.1	.4	15	57/62	134/99	S/N	
006	53.5	—	1.2	.7	18	59/58	120/105	S/N	
007	54.0	59.0	1.3	.7	19	48/52	136/104	S/N	1st subbottom reflector appears midway through segment 6-7.
008	54.5	58.3	2.5	1.3	35	61/70	140/120	S/N	350m of sonar missing. Count adjusted.
009	54.3	—	2.5	.5	—	—	135/115	S/N	no sonar this segment.
010	54.7	—	2.5	.6	22	51/59	140/120	S/N	some BATS missing
011	55.0	—	2.0	1.3	18	46/48	131/111	S/N	
012	54.5	60.0	1.6	1.1	26	40/40	123/103	S/N	2 SUBBOTTOM REFLECTORS.
013	54.7	59.3	1.5	.8	40	32/32	127/107	S/N	
014	53.5	60.0	2.1	1.2	67	40/40	123/103	S/N	
015	52.8	58.0	3.0	1.3	27	41/43	131/111	S/N	upper reflector surfaces at mid segment.
016	53.3	56.5	2.1	1.0	25	44/44	124/104	S/N	gauges become much smaller at this point.
017	53.4	—	1.3	.7	21	39/41	129/109	S/N	2nd sub bottom reflector.
018	51.7	56.0	1.0	.8	22	48/48	127/107	S/N	
019	51.2	54.1	1.0	.5	—	—	—	S/N	no sonar this segment. 3rd Reflector begins this segment.
020	49.7	53.6	2.0	1.1	—	—	—	S/N	2nd reflector surfaces at mid segment.
021	47.3	53.8	1.5	.8	32	72/88	144/109	S/N	
022	47.7	54.0	.6	.6	15	78/85	134/99	S/N	
023	46.3	52.0	1.1	.8	15	85/93	135/100	S/N	
024	45.5	50.4	1.4	.6	20	76/79	132/97	S/N	
025	43.5	49.8	—	—	23	94/102	133/98	S/N	NO FATHOGRAM THIS SEGMENT
026	43.6	—	1.3	1.4	39	93/101	134/99	S/N	3rd Sub bottom reflector surfaces.
027	43.0	—	2.3	.7	34	126/137	135/100	S/N	Large Multi-gauges
028	41.5	—	1.8	1.9	11	114/131	140/105	S/N	

Line Number : 34					Year : 1981				gouge measurements in meters	
Fathogram Measurements					Sonograph Measurements					
SEGMENT	Water Depth (meters)	Reflector "A"	Gouge Depth	Ridge Height	Gouge Width	Density ρ	Orientation (θ T)	Sediment Cohesion	REMARKS	
000	40.2	—	2.0	2.0	15	45/43	76/104	RC		
001	38.0	41.0	3.0	5.0	20	84/81	72/03	RC		
002	35.0	39.0	2.1	2.2	12	138/135	85/95	RC		
003	32.8	37.5	2.0	1.2	8	174/167	73/107	RC	33 m bench	
004	39.0	42.0	1.6	1.2	20	68/65	70/110	RC	muffled gouges?	
005	35.5	41.5	1.6	1.0	12	94/90	78/102	RC	Up over into bench (depression)	
006	32.6	39.0	2.3	1.5	10	90/86	71/109	RC		
007	32.5	36.0	1.8	1.2	10	147/144	86/94	RC		
008	28.5	33.5	2.3	1.5	10	186/182	60/20	RC	2° gouge $\theta = 74^\circ$	
009	27.0	31.5	2.3	1.6	8	169/164	65/115	RC		
010	26.0	30.5	1.0	1.1	5	196/188	72/03	RC		
011	23.5	28.0	1.2	1.1	5	315/306	66/114	RC	SS	
012	20.3	27.0	<.4	<.4	7	216/207	72/108	C	First reflector surfaces, 2nd reflector appears!	
013	19.4	24.2	<.3	<.4	5	25/126	54/126	C	2nd/3rd/4th/5th/6th/7th/8th/9th/10th/11th/12th/13th/14th/15th/16th/17th/18th/19th/20th/21st/22nd/23rd/24th/25th/26th/27th/28th/29th/30th/31st/32nd/33rd/34th/35th/36th/37th/38th/39th/40th/41st/42nd/43rd/44th/45th/46th/47th/48th/49th/50th/51st/52nd/53rd/54th/55th/56th/57th/58th/59th/60th/61st/62nd/63rd/64th/65th/66th/67th/68th/69th/70th/71st/72nd/73rd/74th/75th/76th/77th/78th/79th/80th/81st/82nd/83rd/84th/85th/86th/87th/88th/89th/90th/91st/92nd/93rd/94th/95th/96th/97th/98th/99th/100th/101st/102nd/103rd/104th/105th/106th/107th/108th/109th/110th/111st/112nd/113rd/114th/115th/116th/117th/118th/119th/120th/121st/122nd/123rd/124th/125th/126th/127th/128th/129th/130th/131st/132nd/133rd/134th/135th/136th/137th/138th/139th/140th/141st/142nd/143rd/144th/145th/146th/147th/148th/149th/150th/151st/152nd/153rd/154th/155th/156th/157th/158th/159th/160th/161st/162nd/163rd/164th/165th/166th/167th/168th/169th/170th/171st/172nd/173rd/174th/175th/176th/177th/178th/179th/180th/181st/182nd/183rd/184th/185th/186th/187th/188th/189th/190th/191st/192nd/193rd/194th/195th/196th/197th/198th/199th/200th/201st/202nd/203rd/204th/205th/206th/207th/208th/209th/210th/211st/212nd/213rd/214th/215th/216th/217th/218th/219th/220th/221st/222nd/223rd/224th/225th/226th/227th/228th/229th/230th/231st/232nd/233rd/234th/235th/236th/237th/238th/239th/240th/241st/242nd/243rd/244th/245th/246th/247th/248th/249th/250th/251st/252nd/253rd/254th/255th/256th/257th/258th/259th/260th/261st/262nd/263rd/264th/265th/266th/267th/268th/269th/270th/271st/272nd/273rd/274th/275th/276th/277th/278th/279th/280th/281st/282nd/283rd/284th/285th/286th/287th/288th/289th/290th/291st/292nd/293rd/294th/295th/296th/297th/298th/299th/300th/301st/302nd/303rd/304th/305th/306th/307th/308th/309th/310th/311st/312nd/313rd/314th/315th/316th/317th/318th/319th/320th/321st/322nd/323rd/324th/325th/326th/327th/328th/329th/330th/331st/332nd/333rd/334th/335th/336th/337th/338th/339th/340th/341st/342nd/343rd/344th/345th/346th/347th/348th/349th/350th/351st/352nd/353rd/354th/355th/356th/357th/358th/359th/360th/361st/362nd/363rd/364th/365th/366th/367th/368th/369th/370th/371st/372nd/373rd/374th/375th/376th/377th/378th/379th/380th/381st/382nd/383rd/384th/385th/386th/387th/388th/389th/390th/391st/392nd/393rd/394th/395th/396th/397th/398th/399th/400th/401st/402nd/403rd/404th/405th/406th/407th/408th/409th/410th/411st/412nd/413rd/414th/415th/416th/417th/418th/419th/420th/421st/422nd/423rd/424th/425th/426th/427th/428th/429th/430th/431st/432nd/433rd/434th/435th/436th/437th/438th/439th/440th/441st/442nd/443rd/444th/445th/446th/447th/448th/449th/450th/451st/452nd/453rd/454th/455th/456th/457th/458th/459th/460th/461st/462nd/463rd/464th/465th/466th/467th/468th/469th/470th/471st/472nd/473rd/474th/475th/476th/477th/478th/479th/480th/481st/482nd/483rd/484th/485th/486th/487th/488th/489th/490th/491st/492nd/493rd/494th/495th/496th/497th/498th/499th/500th/501st/502nd/503rd/504th/505th/506th/507th/508th/509th/510th/511st/512nd/513rd/514th/515th/516th/517th/518th/519th/520th/521st/522nd/523rd/524th/525th/526th/527th/528th/529th/530th/531st/532nd/533rd/534th/535th/536th/537th/538th/539th/540th/541st/542nd/543rd/544th/545th/546th/547th/548th/549th/550th/551st/552nd/553rd/554th/555th/556th/557th/558th/559th/560th/561st/562nd/563rd/564th/565th/566th/567th/568th/569th/570th/571st/572nd/573rd/574th/575th/576th/577th/578th/579th/580th/581st/582nd/583rd/584th/585th/586th/587th/588th/589th/590th/591st/592nd/593rd/594th/595th/596th/597th/598th/599th/600th/601st/602nd/603rd/604th/605th/606th/607th/608th/609th/610th/611st/612nd/613rd/614th/615th/616th/617th/618th/619th/620th/621st/622nd/623rd/624th/625th/626th/627th/628th/629th/630th/631st/632nd/633rd/634th/635th/636th/637th/638th/639th/640th/641st/642nd/643rd/644th/645th/646th/647th/648th/649th/650th/651st/652nd/653rd/654th/655th/656th/657th/658th/659th/660th/661st/662nd/663rd/664th/665th/666th/667th/668th/669th/670th/671st/672nd/673rd/674th/675th/676th/677th/678th/679th/680th/681st/682nd/683rd/684th/685th/686th/687th/688th/689th/690th/691st/692nd/693rd/694th/695th/696th/697th/698th/699th/700th/701st/702nd/703rd/704th/705th/706th/707th/708th/709th/710th/711st/712nd/713rd/714th/715th/716th/717th/718th/719th/720th/721st/722nd/723rd/724th/725th/726th/727th/728th/729th/730th/731st/732nd/733rd/734th/735th/736th/737th/738th/739th/740th/741st/742nd/743rd/744th/745th/746th/747th/748th/749th/750th/751st/752nd/753rd/754th/755th/756th/757th/758th/759th/760th/761st/762nd/763rd/764th/765th/766th/767th/768th/769th/770th/771st/772nd/773rd/774th/775th/776th/777th/778th/779th/780th/781st/782nd/783rd/784th/785th/786th/787th/788th/789th/790th/791st/792nd/793rd/794th/795th/796th/797th/798th/799th/800th/801st/802nd/803rd/804th/805th/806th/807th/808th/809th/810th/811st/812nd/813rd/814th/815th/816th/817th/818th/819th/820th/821st/822nd/823rd/824th/825th/826th/827th/828th/829th/830th/831st/832nd/833rd/834th/835th/836th/837th/838th/839th/840th/841st/842nd/843rd/844th/845th/84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[illegible]

[illegible]

Line Number : 36									
Year : 1981									
gouge measurements in meters									
Fathogram Measurements					Sonograph Measurements				
SEGMENT	Water Depth (meters)	Reflector "A"	Gouge Depth	Ridge Height	Gouge Width	Density	Orientation (°T)	Sediment Cohesion	150m range REMARKS
000	5.0	—	<.2	<.2	1	2/3	21/78	S N	funny bottom (still clay) & sand (?)
001	7.3	—	—	—	—	0/0	—	S N	
002	9.2	—	<.2	<.2	1	1/1	14/121	S N	funny bottom
003	12.3	—	<.2	<.2	1	10/11	44/101	S N	
004	13.7	—	.3	.3	4	26/25	75/110	S N	
005	15.6	—	<.3	<.2	4	34/35	50/107	S N	Gauging in subbottom : sand gauging
006	16.9	—	.3	.3	4	25/25	55/112	S C	
007	17.7	—	.2	.2	7	78/79	55/112	S C	
008	18.5	—	.2	.3	5	23/23	62/112	S N	
009	20.1	—	<.2	<.2	5	50/49	62/112	S N	
010	21.2	—	.2	.3	6	80/92	40/39	S N	
011	21.6	—	.2	.4	6	126/137	43/114	R C	
012	22.8	—	.5	.5	10	196/204	52/111	R C	1 m shoal/bench. gauging heavier to seaward
013	26.2	—	1.2	1.1	8	176/172	118/143	R C	
014	—	—	.6	.6	5	109/105	102/127	R C	
015	30.5	33.5	1.0	1.2	5	119/120	123/148	R C	gouges appear SUBSIDED.
016	34.0	37.2	.7	.5	5	92/90	120/145	R C	
017	34.0	39.5	1.0	1.0	8	69/66	105/130	R C	
018	35.5	40.2	.6	.4	6	82/80	118/143	R C	
019	35.6	40.5	.8	1.3	7	102/102	90/115	R C	
020	35.7	39.3	.3	1.0	6	111/108	117/142	R C	
021	36.5	41.0	.6	.8	5	136/131	70/138	R C	
022	38.4	40.0	.8	.5	6	89/86	65/120	R C	gouges appear fresh again.
023	39.5	42.0	1.5	1.4	10	79/76	118/278	R C	undulating silt b
024	37.1	39.1	1.4	.9	10	99/100	125/285	R C	
025	35.6	39.7	.4	.5	5	151/183	125/285	R C	
026	34.5	39.0	.7	1.6	8	186/188	125/285	R C	
027	34.4	38.0	.8	.9	8	165/190	138/298	R C	
028	34.2	40.0	.7	.6	12	164/179	135/298	R C	

Line number : 36 (cont.)									
Year : 1981									
gauge measurements in meters									
M									
	width	height	depth	width	depth	width	depth	width	depth
029	11.0	0.0	0.0	10	118/117	124/1284	RC	150 m	REMARKS
030	34.2	0.0	0.0	10	121/120	124/1284	RC	150 m	REMARKS
031	11.0	0.0	0.0	10	121/120	124/1284	RC	150 m	REMARKS
032	11.0	0.0	0.0	10	121/120	124/1284	RC	150 m	REMARKS
033	22.4	0.0	0.0	5	175/177	123/283	RC	150 m	REMARKS
34	31.6	0.0	0.0	7	190/220	124/1284	RC	150 m	REMARKS
36	29.5	0.0	0.0	0	148/452	140/1500	RC	150 m	REMARKS
37	27.0	0.0	0.0	12	235/244	130/290	RC	150 m	REMARKS
38	27.0	0.0	0.0	12	510/501	115/1275	RC	150 m	REMARKS
39	27.0	0.0	0.0	10	120/121	124/1284	RC	150 m	REMARKS
40	24.7	0.0	0.0	10	204/198	140/1180	RC	150 m	REMARKS
41	24.3	0.0	0.0	10	100/100	124/1284	RC	150 m	REMARKS
42	22.2	0.0	0.0	5	169/164	115/287	RC	150 m	REMARKS
43	21.5	0.0	0.0	1	124/125	120/1247	RC	150 m	REMARKS
44	11.0	0.0	0.0	0	125/132	122/1202	RC	150 m	REMARKS
45	10.8	0.0	0.0	3	53/54	127/297	SN	150 m	REMARKS
46	17.5	0.0	0.0	4	33/36	135/300	SN	150 m	REMARKS
47	17.0	0.0	0.0	6	64/62	142/297	SN	150 m	REMARKS
48	11.0	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
49	12.5	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
50	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
51	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
52	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
53	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
54	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
55	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
56	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
57	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
58	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
59	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
60	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
61	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
62	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
63	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
64	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
65	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
66	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
67	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
68	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
69	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
70	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
71	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
72	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
73	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
74	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
75	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
76	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
77	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
78	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
79	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
80	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
81	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
82	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
83	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
84	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
85	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
86	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
87	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
88	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
89	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
90	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
91	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
92	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
93	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
94	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
95	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
96	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
97	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
98	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
99	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
100	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
101	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
102	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
103	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
104	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
105	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
106	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
107	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
108	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
109	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
110	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
111	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
112	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
113	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
114	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
115	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
116	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
117	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
118	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
119	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
120	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
121	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
122	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
123	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
124	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
125	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
126	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
127	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
128	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
129	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
130	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
131	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
132	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
133	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
134	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
135	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
136	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
137	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
138	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
139	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
140	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
141	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
142	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
143	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
144	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
145	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
146	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
147	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
148	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
149	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
150	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
151	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
152	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
153	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
154	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
155	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
156	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
157	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
158	11.4	0.0	0.0	7	124/125	120/1247	RC	150 m	REMARKS
159									

Line number : 37

Year : 1981

gouge measurements in meters

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